February 2013

## **REPORT ON**

# Groundwater Vulnerability Study Richmond Village Well System Richmond, Ontario

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REPORT

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## **Executive Summary**

The Village of Richmond (Village) is located approximately 35 km southwest of downtown Ottawa, along the Jock River. Water demand from the Village is currently met by a combination of individual private wells and two communal supply systems, referred to herein as the King's Park (municipal) and Hyde Park (private) systems. Groundwater vulnerability assessments were completed previously for these existing well systems, which involved construction and calibration of a groundwater flow model that was used to define the Well Head Protection Areas (WHPAs) associated with these well systems (Golder, 2009 and 2010a).

In 2008, Mattamy Homes Limited (Mattamy) initiated a Water and Sewer Master Servicing Study for the Village of Richmond, which recommended communal groundwater supply wells as the preferred option for new development on the lands located on the western portions of the existing Village (Stantec, 2011). A drilling and testing program was subsequently undertaken, where a number of wells were completed in the Nepean Formation aquifer. Portions of the development lands were acquired in 2010 by Richmond Village (South) Limited and Richmond Village (North) Limited, together referred to as RV. Golder Associates Ltd. (Golder) was retained by RV to complete a groundwater vulnerability study for the proposed municipal wells PW08-1, PW09-1 and PW09-2.

The groundwater flow model previously used in the King's Park and Hyde Park studies was used as the starting point for development of time-of-travel capture zones for the RV wells. The model parameterization was found to be consistent with data collected subsequent to its original development (i.e. hydraulic testing of the RV wells indicated an aquifer transmissivity that fell within the range used in the previous studies). Adjustments to the finite difference grid were required in order to properly include the RV wells. The pumping rate used in the capture zone delineation was based on the calculated average water demand for the Richmond development at full build-out (1,630  $m^3/d$ ) and was assumed to be shared evenly between PW09-1 and PW08-1.

Groundwater vulnerability mapping was performed over the area of the model domain using the Intrinsic Susceptibility Index (ISI) method, and intrinsic vulnerability scores were calculated based on the results of the groundwater vulnerability mapping and WHPA zone delineation. A threats assessment for the existing King's Park and Munster wells was previously undertaken by Dillon Consulting on behalf of the Mississippi-Rideau Source Protection Region (MRSPR), and was used as the basis for the threats assessment for the RV wells. In accordance with direction from the Source Protection Office of the Ministry of the Environment, only threats that could be considered significant according to the established methodology require evaluation.

WHPAs for the RV wells were similar to those developed previously for the King's Park wells; the capture zones extended in two directions, extending approximately 14 km towards the west-northwest, and approximately 6 km towards the south. Following the application of the ISI method, an aquifer vulnerability score of "low" was defined throughout the WHPA zones, which resulted from the extensive thickness of overlying geological materials found in the study area. Calculated vulnerability score ranged from 10 (limited to WHPA Zone A), to 2 in WHPA Zone D.

The uncertainty associated with the vulnerability scores within the WHPA was determined in accordance with Draft Guidance Module 3; the areas encompassed within WHPA Zone B were categorized as "low uncertainty", where the remaining areas within the overall WHPA were categorized as "high uncertainty".



Based on the results of the vulnerability scoring, significant threats to the RV wells, exclusive of DNAPL's, can only occur within WHPA Zone A. The current land use on the site is agricultural, but this will change to residential/parkland as development proceeds. Based on the current methodology, sewage connections and laterals are defined as wastewater collection facilities and are considered significant threats in areas, such as WHPA Zone A, with an intrinsic vulnerability score of 10. The draft Source Protection Plan prepared by the MRSPR includes Policies SEW-6-LB and SEW-7-LB-PI-MC. The former includes requirements for the inspection and maintenance of sanitary sewers and related pipes where they are, or would be considered a significant drinking water threat. Policy SEW-7-LB-PI-MC includes a requirement that approval under the Ontario Water Resources Act includes appropriate measures to manage the threat. This includes minimum construction standards for new or replacement sanitary sewers and related pipes that can be required by the director of the MOE as appropriate.

The storage and handling of DNAPL's is considered a significant threat within WHPA Zones A, B and C (within a 5 year time-of-travel). The current WHPAs were reviewed and compared to those of the Richmond King's Park assessments. Both computer and on-ground surveys were used to verify that no additional sources, such as dry cleaners, manufacturing facilities or wood product manufacturers are found within the relevant WHPAs.

Significant data gaps identified in this study relate to the characterization of the porosity of the hydrostratigraphic units and the bedrock fracture network. Additional data relating to these parameters would provide a better estimate of the groundwater travel times and could be used to improve the delineation of the WHPAs. Additional data gaps for the current study relate to the threats identified within areas of overlap between the current WHPA Zone C and those of the King's Park assessment. These areas should be re-assessed during the next source protection iteration. Additionally, although the best data available at the time of this assessment were used to define the rates of groundwater pumping from other groundwater users found within the study area, these data could be refined to improve the overall WHPA delineation.

The groundwater vulnerability study has confirmed that risk to the proposed new supply wells is minimal due to the low vulnerability of the supply aquifer and the absence of significant threats to the water supply. It is not anticipated that extraordinary aquifer protection measures will be required for the proposed system.





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## 1.0 INTRODUCTION

## 1.1 Background

The Village of Richmond (Village) is located approximately 35 km southwest of downtown Ottawa, along the Jock River (Figure 1). Water demand from the Village is currently met by a combination of individual private wells and two communal supply systems, referred to herein as the King's Park, and Hyde Park systems. King's Park is a municipally owned and operated system, while Hyde Park is privately owned and operated. The King's Park and Hyde Park systems are currently comprised of two supply wells each, and service approximately 151 and 94 properties respectively within the Village (Stantec, 2011, Golder 2010a). A third Hyde Park well was constructed in 2010 and will be incorporated into the system to meet increasing system demands. Waste water within the Village is provided by the central municipal sanitary collection and treatment system.

The Clean Water Act 2006 (CWA) requires that vulnerability studies be prepared for all municipal drinking water systems that use groundwater. A vulnerability assessment of the King's Park communal well system was completed by Golder Associates Ltd. (Golder) in 2003 as a part of a study for the Mississippi-Rideau Source Protection Region (MRSPR), and subsequently updated in 2008 and 2009 (Golder, 2003, 2008b and 2009). This work involved the construction and calibration of a groundwater flow model, which was used to estimate the wellhead protection areas (WHPAs) for the wells. While the Hyde Park development is private and therefore exempt from the requirements of the CWA, the City of Ottawa requires that wellhead protection studies similar to those required under the CWA be undertaken for private communal well systems. The existing model was used to develop the WHPAs associated with the Hyde Park wells (Golder, 2010a).

In 2008, Mattamy Homes Limited (Mattamy) initiated a Water and Sewer Master Servicing Study (MSS) for lands located on the western portions of the Village, which recommended groundwater supply wells as the preferred option for new development on these lands (Stantec, 2011). Groundwater wells were subsequently installed and tested at the locations shown on Figure 2. Portions of the development lands were purchased in 2010 by Richmond Village (South) Limited and Richmond Village (North) Limited, together referred to as RV in this report. These lands included the wells PW09-1, PW09-2, and PW08-1. Golder was retained by RV to complete a groundwater vulnerability study for these wells.

## 1.2 Scope of Work

The technical requirements, study approach and methodology for groundwater vulnerability studies are set out within the Ontario Clean Water Act, 2006, Technical Rules: Assessment Report (November, 2009), as well as the draft Guidance Module 3 (October, 2006) prepared by the MOE in advance of the development of regulations, rules and final guidance documents under the Clean Water Act.

The objective of the study was to identify wellhead protection areas (WHPAs) for the proposed communal well system, to map the relative vulnerability of the supply aquifer within the identified WHPAs as high, medium or low, and to assign intrinsic vulnerability scores within the WHPAs.

The scope of work included:

- Reviewing the existing conceptual and numerical groundwater flow models that were developed as a part of the 2003 wellhead protection study;
- Reviewing hydrogeological data collected subsequent to completion of the 2003 study, including aquifer characteristics (thickness, transmissivity and storage properties) and additional nearby permitted water takings;



- Updating the groundwater flow model as necessary based on the review of data noted above;
- Using results of the recalibrated numerical groundwater flow model to assist in delineation of the WHPA zones A, B, C, and D using the forecasted pumping rates as defined in the MOE Guidance Module 3;
- Assessing the vulnerability of the supply aquifer based on a calculation of the Intrinsic Susceptibility Index (ISI), using the available hydrogeological data and the results of the numerical groundwater modelling;
- Using the results of the ISI calculation to categorize the relative vulnerability within the WHPAs as high, medium or low using Table 4.1 in Draft Guidance Module 3;
- Assigning intrinsic vulnerability scores within the WHPAs based on the delineated WHPA zones and results
  of the ISI vulnerability assessment,
- Completing an assessment of the uncertainty of the vulnerability scoring and categorizing the uncertainty as either high or low; and,
- Completing an external peer review to evaluate the suitability of the scope, methodology, and results of the current study. The details of the peer review, prepared by Dillon Consulting are included in Appendix A.

## 1.3 Proposed Water Supply System

The Mattamy Richmond well system is comprised of three wells, referred to as PW08-1, PW09-1, and PW09-2. PW08-1 was constructed in October 2009 and was completed at a depth of 137.16 m at the location shown on Figure 2. This well is completed with casing installed to a depth of 45.72 mbgs, and is open below this depth. In December 2009, a 0.254 m diameter production well (PW09-1) was drilled to a depth of 70 mbgs. PW09-1 is located near the eastern property boundary, approximately 650 m south of Perth St (Figure 2). The production well was completed with 45.72 m of steel well casing that was grouted in place. The steel casing was installed through the upper portion of the Oxford Formation, and groundwater flow to the well is expected to occur primarily from the lower aquifer (a description of the hydrostratigraphy of the site is provided in Section 2.4). In January, 2010, a third well (PW09-2, 0.254 m diameter) was installed less than 5 m from PW09-1. Similar to well PW09-1, PW09-2 is completed to a depth of 77.72 mbgs, with steel well casing installed to a depth of 45.72 mbgs. Further details of the Richmond wells are provided in Table 1 and well records are provided in Appendix B.

Forecast water demands on the Richmond well system were calculated based on a projection of the population growth for the development. The number of planned housing units and assumed water demand requirements for the Richmond development was provided by representatives of Mattamy and RV, summarized as follows:

- The RV lands will contain 1,000 units, including 650 single homes and 350 town homes;
- The Mattamy lands will contain 1,000 single homes;
- The average water demand is 835 L/d/unit for single homes and 720 L/d/unit for town homes (Stantec, 2011); and,

Therefore, the anticipated average water demand for the Richmond development at full build-out is  $1,630 \text{ m}^3/\text{d}$  (1,132 L/min).



## 2.0 REGIONAL SETTING

## 2.1 Physiography

Figure 1 illustrates the extent of the study area, which covers approximately 880 km<sup>2</sup>, extending from Mississippi Lake in the west to the Rideau River in the east. A number of smaller creeks, streams, ponds, and tributaries to the larger rivers occur throughout the area, including the Jock River, a tributary to the Rideau River that flows through the Village of Richmond.

Topography within the study area generally slopes towards the east (Rideau River), with elevations ranging from approximately 145 meters above sea level (masl) along the western portions of the study area to approximately 85 masl along the eastern portions of the study area. Ground elevations are approximately 90 masl within the Village. Local topographic variations beyond the above noted range occur in isolated areas within the overall study area.

## 2.2 Surficial Geology

The surficial geology of the study area is illustrated on Figure 3. Surficial deposits found throughout the study area consist of glacial till, sandy and coarser stratified ice contact sediments, sandy and coarser nearshore sediments, a dense marine clay unit and modern fluvial deposits. The glacial till unit outcrops mainly in the western part of the study area where its thickness varies from centimetres to several metres. Many of the till outcrop areas are covered by large (several km<sup>2</sup>) marshes and wetlands. An ice contact stratified deposit is reported 5 km northeast of Richmond. In the eastern part of the study area (including Richmond) these units are overlain by a marine clay material which increases in thickness towards the east. In the Richmond area, the overburden deposits mainly consist of the marine clay, underlain by the glacial till unit. Modern fluvial deposits are present along the Jock River.

## 2.3 Bedrock Geology

The bedrock geology of the study area is illustrated on Figure 4. In general, the surficial deposits within the study area are underlain by sedimentary rocks of Paleozoic age composed of sandstones, dolostones, limestones and shales. This sedimentary sequence unconformably overlies the Precambrian basement which underlies all of the study area and only outcrops west of it (near Mississippi Lake).

The Precambrian basement consists of highly deformed metasedimentary rocks. The Precambrian basement is overlain by the Nepean Formation sandstones which outcrop in the western area of the study, along the eastern shore of Mississippi Lake. The Nepean Formation, which is the upper unit of the Potsdam Group, underlies all of the study area. The lower member of the Potsdam Group, the Covey Hill Formation, is not represented in the study area. The Nepean Formation is composed of alternating beds of calcareous sandstone and quartz arenite. Its thickness, as reported by previous authors and estimated in the scope of the present study, would reach 45 m in the Richmond area. It is considered the most transmissive aquifer within the study area. Flow in this aquifer is controlled predominantly by fractures, as the primary porosity of the sandstones has been reduced by cementation. The Mattamy Richmond wells are completed in this formation.

Except where it outcrops, the Nepean Formation is overlain by the March and Oxford Formations of the Beekmantown Group. The March Formation consists of interbedded quartz sandstone, sandy dolostone/dolostone, and shale partings are present through the formation. Thicknesses up to 6.6 m have been documented for the March Formation in the study area. The Oxford Formation is mainly composed of dolostone with commonly



occurring shale interbeds, and thicknesses up to 60 m have been reported in the study area. Both the March and Oxford Formations are considered as good carbonate rock type aquifers, though their capacities are lesser than that of the Nepean Formation.

The Rockcliffe Formation, of the Chazy Group, is found mostly in the western part of the modelled area. It consists mainly of interbedded quartz sandstone and shale. Previous authors have estimated its thickness to be in the order of 10 m. The Rockcliffe Formation is typically not considered a significant water producing aquifer.

The Gull River and Bobcaygeon Formations, of the Ottawa Group, are also found mostly in the western part of the study area. The Gull River Formation is made of interbedded limestone and silty dolostone at the bottom and of finely crystalline limestone at the top. The Bobcaygeon Formation is described as an interbedded lithographic to coarsely crystalline limestone with shale parting. The Ottawa Group formations are referred to as marginal and variable water producing aquifers, although often able to provide enough water for an individual residence.

The March, Oxford, Rockcliffe, Gull River and Bobcaygeon outcrops are numerous and cover large areas west of Richmond. Several faults and joint sets have been mapped in these units. The faults in this area form part of a major tectonic feature, the Ottawa-Bonnechere Graben. Faults are steeply dipping and generally strike in an east to southeast direction. Vertical displacements reported along these faults generally vary between 5 and 50 m in the study area. Specific information on the hydraulic characteristics of the faults in the bedrock was not available local to the Richmond wells. It is not known whether the faults represent barriers or conduits for groundwater flow in this area. However, available information pertaining to the dissolved contaminant plume in the Blacks Corners area (approximately 25 km west of Richmond) suggests that groundwater flow occurs across faults in this area. The Paleozoic rocks are also intersected by steeply dipping (near vertical) joints which form four sets (N015°, N055°, N100° and N145°). Joint spacing typically varies from 30 to 100 cm and the N100° joint set is reported to be the most dominant (Golder, 2003).

## 2.4 Hydrogeology

## 2.4.1 Hydrostratigraphy

There are two primary bedrock aquifer systems that are identified within the study area. The upper aquifer is typically defined within the upper 35 m of the Oxford Formation. The majority of the private residential wells within Richmond are completed within and obtain water from this aquifer. The lower aquifer is defined within the lower portion of the March formation and the upper portion of the Nepean Formation. In general, the lower aquifer is considered to be more transmissive compared to the upper aquifer. Currently operating communal wells in the area (the King's Park and Hyde Park wells in Richmond, in addition to wells in Almonte, Munster, Kemptville, and Merrickville) draw water from the lower aquifer. In some areas the two aquifers are separated by a bedrock aquitard consisting of limestone of the lower Oxford Formation and interbedded limestone and sandstone of the March Formation. The presence of this aquitard is often indicated by strong upward vertical gradients between the aquifers. The potentiometric surface of the lower aquifer is typically above ground surface, and wells completed in this aquifer often flow.

The overburden within the study area consists of glacial till and marine clay deposits which are not considered suitable for water supply (only one of 893 water supply wells in the Village of Richmond was completed in the overburden (Golder, 2008a)). The Nepean formation is underlain by highly deformed Precambrian-aged metasedimentary rock, which is generally used as a supply aquifer only where neither the Paleozoic bedrock aquifers nor overburden aquifers are present.





#### 2.4.2 Groundwater Flow

Observed groundwater flow characteristics within the study area are shown on Figure 5. Groundwater within the study area is generally interpreted to flow from the Mississippi River and Mississippi Lake in the west towards the Rideau River in the east. Upward gradients from the Nepean to the Oxford formation have been observed in Black's Corners and Manotick, as wells completed at the Site (Golder, 2011). Upward gradients are also expected where the Jock River flows over outcrops of the Oxford formation to the south of Munster (Golder, 2003). Shallow groundwater flow in the overburden units is expected to vary as a function of bedrock topography (Golder, 2010a).

#### 2.4.3 Groundwater Quality

The lower sandstone aquifer (Nepean Formation) underlying the site is regionally extensive, and is utilized by the King's Park and Hyde Park communal wells in Richmond, as wells as the communal wells systems in Almonte, Munster, Kemptville and Merrickville. The groundwater in the lower aquifer is hard (typical for groundwater sources, particularly those that have a substantial carbonate composition) and occasionally exceeds the non-health related aesthetic criteria for iron. The exceedances of the aesthetic criteria for iron in the lower aquifer are generally treatable using conventional water softening. Where both aquifers are present, the shallower Oxford formation is typically used for residential supply purposes and tends to be of somewhat lower quality than the deeper aquifer. The groundwater in the upper aquifer is quite hard and occasionally exceeds the aesthetic criterion for iron. Low concentrations of hydrogen sulphide are often present. Overall, wells completed in either supply aquifer are expected to produce groundwater that is safe and aesthetically suitable for human consumption (Golder, 2008a).





## 3.0 SUMMARY OF PREVIOUS STUDIES

## 3.1 Hydrogeological Evaluations

Hydraulic testing data were collected for the overburden, Oxford, March, and Nepean bedrock formations as a part of previous hydrogeological investigations completed by Golder and others within the study area. These data are summarized in Table 2. The following describes the results of this testing. Tests completed in the Oxford Formation (Golder 2006), which are included in the Table for completeness, were not included in the following discussion, as the focus of the current study is on the Nepean Formation.

The hydraulic conductivity of the silty clay and fine sand overburden materials was estimated based on hydraulic response testing (rising head tests) of monitoring wells located on the Richmond property to range between approximately  $1 \times 10^{-6}$  m/s and  $1 \times 10^{-5}$  m/s (Golder, 2010b). This range is representative of the conditions local to the site, which is not necessarily representative of conditions across the full study area.

As a component of the Wellhead Protection Study for the King's Park wells completed by Golder in 2003, a 6-hour pumping test was completed on King's Park well KP2 on July 10, 2002. The well was completed as an open hole through the Oxford and March formations, and into the Nepean formation. It is interpreted that the majority of the water supplying the communal wells comes from the Nepean formation. The well was pumped at a rate of 1,100 m<sup>3</sup>/day for the 6-hour period. Based on the results of the pumping test, the transmissivity and storage coefficient of the aquifer were estimated to be 605 m<sup>2</sup>/day and 1.9 x 10<sup>-5</sup>, respectively (Golder, 2008a).

Pumping tests were completed on three wells in the Hyde Park area of Richmond Village as a component of various hydrogeological studies completed for that development (Golder, 2010a). The wells varied in depth from 83 to 94 m, and were completed as open holes throughout the Nepean, March, and Oxford formations. Each well was pumped for a period of 24-hours. Analysis of these tests resulted in a range of transmissivity from 21 to 130 m<sup>2</sup>/day. It is noted that although these wells were completed across several hydrostratigraphic units it is assumed that the majority of the flow was derived from the Nepean formation.

In November 2009, a 48-hour pumping test was conducted on PW08-1 (located within Richmond Village, see Figure 2, and well description in Section 1.3) using a pumping rate of 1,273 Litres per minute (L/min). The transmissivity and storativity values generated by the analysis of drawdown data from the pumping test range from 328 metres squared per day ( $m^2/day$ ) to 700  $m^2/day$  and from 9 x 10<sup>-4</sup> to 1 x 10<sup>-2</sup>, respectively. Based on the results of the pumping test, the sustainable yield of the well was estimated to be 2,600 L/min (Golder, 2011).

A 72-hour pumping test was conducted on well PW09-1 (located within Richmond Village, see Figure 2, and well description in Section 1.3) between September 27 and September 30, 2011. Recovery measurements were collected until October 3, 2011. The pumping test was started at a rate of 2,690 L/min. After the first day of pumping, the rate decreased slightly due to a loss of pump efficiency. The remainder of the test was completed at a rate that ranged from 2,690 L/min to 2,410 L/min. During the pumping test, water level data was collected from the pumping well (PW09-1) and four observation wells screened within the Nepean formation (PW09-2, PW08-1, MW08-1A, MW08-1B), as well as five additional wells screened in the upper aquifer. Based on analysis of pumping and recovery data the transmissivity of the lower sandstone aquifer was estimated to range from 500 m<sup>2</sup>/day to 800 m<sup>2</sup>/day. The pumping test results indicated that the sustainable yield for well PW09-1 is at least the minimum pumping rate of 2,410 L/min and is likely greater.





Hydraulic conductivity of the Precambrian unit underlying the Nepean formations is estimated to range from  $1 \times 10^{-8}$  to  $1 \times 10^{-7}$  m/s based on values reported in surrounding areas (Golder, 2003). Information on the hydraulic characteristics of the Precambrian unit within the study area was not available.

## 3.2 Groundwater Vulnerability Studies

A Wellhead Protection Study (WHPS) was completed by Golder for the communal wells in the King's Park Subdivision in April 2003 (Golder, 2003). A 3-dimensionsal numerical model (MODFLOW) was developed for the study area, and a modeling exercise was completed to define the time-related groundwater capture zones for the King's Park wells. Capture zones were determined using MODPATH to release particles over the open intervals of the wells–, for the "base-case" model and for 21 additional modelling scenarios employed in a sensitivity analysis. The capture zones were therefore "composite" areas, combining the particle tracking results for the 22 modelling scenarios. Due to the significant depth of the lower aquifer below ground surface and the nature of the overlying bedrock formations, the aquifer vulnerability was classified as low throughout the Munster and King's Park WHPAs (Golder, 2003). A recommended well-head protection strategy was proposed as a part of that study.

In October 2006, a series of draft guidance modules were provided by the MOE as part of the *Clean Water Act*. Draft Module 3 – Groundwater Vulnerability Analysis provided new technical requirements and methodologies for defining WHPAs (MOE, 2006). In May 2008, Golder produced 5-year time of travel (ToT) capture zones for the Munster and Kings Park municipal wells using the previously constructed MODFLOW model (Golder, 2008b). The capture zones were determined using the same methodology followed in 2003.

In May 2009 capture zone modelling and an aquifer vulnerability assessment (Golder, 2009) was completed for the upper (Oxford Formation) aquifer using the previously constructed Munster-Kings Park model. For that study, the model was revised in order to relocate the municipal wells (using co-ordinates provided by MRSPR) and to subdivide the upper aquifer layer into three layers. A vulnerability assessment for the upper aquifer was then completed using Intrinsic Susceptibility Index approach (ISI).

In 2010 the Munster-Kings Park Model was used to develop a Well Head Protection Area Plan (WHPAP) for future phases of the Hyde Park development in Richmond (Golder, 2010a). The pumping tests discussed above were a component of that study. Based on the results of that study, the modelled supply rate for the Hyde Park well was increased to  $265 \text{ m}^3$ /day.





## 4.0 STUDY METHODOLOGY

## 4.1 Groundwater Flow Modelling

A groundwater flow model encompassing the study area was constructed previously as a part of the 2003 groundwater vulnerability study for the Munster and King's Park well systems (Golder, 2003). Upon review of the hydrogeological data that has become available since the completion of the 2003 study (described in Section 3.0 above), including data from borehole drilling, aquifer testing, and groundwater level monitoring, it was determined that these data are generally consistent with the original hydrogeological conceptualization. Therefore, the original groundwater flow model developed for the 2003 study is considered appropriate for use in the current study.

The groundwater flow model is described in the sections below. In general, this text follows that which was included in the 2003 assessment report to describe the groundwater flow model, though additional discussion is included herein to document the minor changes that were made to the model to incorporate the Richmond pumping wells.

The overall objective of the groundwater flow modelling was to delineate time-of-travel capture zones for the Richmond supply wells based on forecast pumping rates. These capture zones are subsequently used in the generation of WHPAs for the wells, forming the basis for the groundwater vulnerability assessment.

#### 4.1.1 Modelling Scope

A numerical groundwater flow model was used to assess the time of travel capture zones, as per the *Clean Water Act, 2006* Technical Rules: Assessment Report (November 16, 2009), as well as other applicable MOE guidance, such as the draft Guidance Module: Groundwater Vulnerability Analysis (MOE, 2009 and 2006). Specifically, to meet the above objectives, the following tasks were completed:

- Review of geological and hydrogeological data that has become available since the time of the original groundwater vulnerability assessment;
- Determine the suitability of the above-noted data within the context of the existing conceptual model;
- Update the existing groundwater flow model to account for potential refinement of the conceptual model (as required);
- Verify the location and forecasted water usage for the Richmond Wells, and incorporate these wells into the existing groundwater flow model;
- Verify the calibration of the groundwater flow model using available water elevation data (MOE water well data, and site-specific observation well data), including data collected since the time of the original assessment; and,
- Delineate capture zones using forecasted pumping rates.

An additional task was added as a result of requests from the peer reviewer and the MRSPR. The capture zones for the King's Park wells and the Munster Hamlet wells were updated to reflect anticipated changes due to the operation of the new Richmond wells.





#### 4.1.2 Modelling Approach

The objective of the groundwater modelling for the Richmond area was the determination of time-related capture zones for the groundwater supply wells. As per the MOE guidance, the time-related capture zones of interest include the zero to 2 year time of travel (ToT), the 2 to 5 year ToT; and the 5 to 25 year ToT.

The MOE (under the Clean Water Act; MOE, 2009) requires subdivision of the WHPAs into four zones as follows:

- 1) Area WHPA-A, the surface and subsurface area centred on the well with an outer boundary identified by a radius of 100 m;
- 2) Area WHPA-B, the surface and subsurface areas within which the time of travel to the well is less than or equal to two years but excluding WHPA-A;
- 3) Area WHPA-C, the surface and subsurface areas within which the time of travel to the well is less than or equal to five years but greater than two years; and,
- 4) Area WHPA-D, the surface and subsurface areas within which the time of travel to the well is less than or equal to twenty-five years but greater than five years.

A 3D numerical (MODFLOW) groundwater model was constructed and calibrated with available hydrogeological data to estimate the time-related capture zones for the King's Park and Munster water supply wells (Golder, 2003). This model was subsequently adapted in order to estimate the time-related capture zones for the Richmond water supply wells. Data used to develop the model included information from the MOE Water Well Information System (MOE WWIS) as well as information from geological and hydrogeological investigations completed within the study area (inclusive of studies at Beckwith/Blacks Corners, and Manotick). MOE WWIS data from 2003 through 2011 were incorporated into the groundwater flow model to check for reasonableness of calibration using the most recent available data. Similarly, information from hydrogeological investigations completed subsequent to the 2003 assessment (summarized in Section 3.0 of this report) was evaluated to ensure a reasonable fit was maintained between these data and the existing conceptual model.

Following construction and calibration of the groundwater model under current conditions, predictive computer simulations were completed using the forecasted pumping rates to delineate capture zones. The forecasted pumping rates reflect the future average rates based on the total water demands projected for the supply wells at full buildout of the Mattamy/RV lands.

To delineate the capture zones under forecasted rates, groundwater particles were simulated in the numerical model at the pumping wells, backward-tracked using MODPATH, and the resulting particle traces were projected in plan view to ground surface. The time-related capture zones that are subsequently derived from this analysis represent a two-dimensional (2D) projection of the particle outlines to ground surface. The capture zones for the King's Park and Munster well systems were also delineated, as the new Richmond wells will affect them to some degree.

Both the Munster and King's Park well systems have two sets of capture zones, as the wells in both locations are open to both the shallow and deep bedrock aquifers (Golder, 2009). The casing in the new Richmond wells has been extended to prevent any inflow to the wellbore from the upper aquifer, and therefore only one set of capture zones (deep aquifer) have been mapped for these wells.





#### 4.1.3 Code Selection and Description

MODFLOW-2000 (USGS, 2000) and MODPATH were used to estimate the time-related capture zones. MODFLOW is a multi-purpose three dimensional groundwater flow code developed by the United States Geological Survey. It is modular in nature and uses the finite difference formulation of the groundwater flow equation in its solution. Visual MODFLOW® (Version 4.3.0.154) was used as the numeric flow engine for the simulations presented in this report. MODPATH (Pollock, 1989), a companion code to MODFLOW, was used to complete the particle tracking analyses necessary for the capture zone delineation.

#### 4.1.4 General Modelling Assumptions

The use of the MODFLOW/MODPATH groundwater model infers that the groundwater flow system in the Richmond area can be simulated as an "equivalent porous media" at the scale of the time-related capture zones under consideration (i.e., 0 to 2-year ToT; 2 to 5-year ToT; and 5 to 25-year ToT). Under this assumption, the rate of groundwater flow towards a pumping well occurs as a function of the hydraulic gradient, the hydraulic conductivity, and the porosity of the aquifer. While groundwater flow in bedrock aquifers is controlled primarily by fractures, an equivalent porous media approach is usually used to represent groundwater flow in these aquifer systems. This is considered reasonable provided the scale of the observation (i.e., in this case the extent of the capture zone) is much greater than the scale of the individual fractures, and consideration is given to the selection of a reasonable effective porosity for the bedrock.

Modelling calculations were completed at "steady-state". Under this assumption, the predictions reflect longterm average conditions. Therefore, the potential effects of short-term conditions (floods, severe rainfalls, etc.), seasonal fluctuations and long-term climate changes were not considered.

Table 3 provides a summary of the general modelling assumptions used in the groundwater flow model.

#### 4.1.5 Conceptual Model

The conceptual model developed for the 2003 assessment was reviewed in light of new data collected since the previous study. The major aspects of the conceptual model remain unchanged. These include the topography and drainage conditions across the study area, hydrostratigraphic layering, geological material properties, groundwater and surface water flow directions and magnitudes, assumed flow boundaries. Based on the review it was determined that the conceptual model used previously was appropriate as the basis for the current assessment. The reader is referred to the 2003 report for a detailed description of the development of the conceptual model.

#### 4.1.6 Grid Discretization

The model domain is horizontally discretized into 127 rows and 178 columns per layer. Horizontal grid spacing is specified as 500 m near the outer edges of the model domain, transitioning to 5 m in the vicinity of the pumping wells. It should be noted that horizontal grid refinement was required to accommodate the inclusion of the Richmond wells within the existing groundwater flow model. The total number of cells within the model domain is 158,242, inclusive of all 7 numerical layers, as discussed below.

#### 4.1.7 Model Layering and Surfaces

Figure 6 shows the model layering and hydrostratigraphic conceptualization. The groundwater flow model was constructed using one overburden layer and six bedrock layers, defined as follows:





- Overburden (Model Layer 1) The overburden material found within the study area, which is mainly composed of glacial till (overlain by clayey material in the eastern part of the model area) was conceptualized as a "moderate-low" permeability layer of variable thickness. The upper surface of this unit was defined by topographic data (25 m digital elevation model);
- Upper Bedrock (Model Layer 2) The upper bedrock unit, which includes the Paleozoic dolomites, limestones, and shales of the March, Oxford, Rockcliffe, Gull River, and Bobcaygeon Formations, was conceptualized as a "moderate-low" permeability upper bedrock layer of variable thickness. The surface defining the top of the bedrock was constructed using the bedrock depth information from the MOE WWIS and other available borehole data, and subtracting this depth from the ground surface across the study area. Outliers in the data were identified by plotting the residual elevation (i.e. the difference of the input bedrock surface elevations and the interpolated results). Points having a residual elevation greater than 10 m were removed. The interpolated bedrock surface was corrected to the ground surface elevation where the bedrock surface was interpolated to be above ground surface. The resulting bedrock surface elevation and overburden isopach maps are shown on Figure 7 and Figure 8, respectively;
- Nepean Aquifer (Model Layers 3 through 5) Below the upper bedrock layer lies a relatively high to very high conductivity bedrock layer representing the Nepean aquifer. The Nepean Formation surface was constructed as follows: a preliminary surface was constructed using only the wells, boreholes, and bedrock outcrops having clearly identified the Nepean top contact. The MOE WWIS data was filtered to retain only the wells presenting a sandstone bedrock contact elevation at a logical location in the stratigraphic sequence (e.g. any well having limestone or sand and gravel below the sandstone contact depth was removed). The MOE WWIS data were compared against the preliminary surface, and any point from the MOE WWIS dataset having a residual elevation of 20 m or greater was removed. A final surface (illustrated on Figure 9) and thickness of the overlying "upper bedrock" (illustrated on Figure 10) was created using the filtered MOE WWIS data combined with the data used to create the preliminary surface. In order to better match the available data, the Nepean aquifer was discretized into three separate layers, where:
  - The top Nepean layer (model layer 3) is one metre thick and represents fractured sandstones of relatively high hydraulic conductivity. This layer has the same hydraulic properties (hydraulic conductivity, storativity, and porosity) throughout the model area;
  - The middle Nepean layer (model layer 4) represents, depending on the location, either fractured sandstones of relatively very high hydraulic conductivity (same as model layer 3) or sandstones of relatively high hydraulic conductivity. This layer has a constant thickness of 4 m through all of the model area, but its conductivity is lower in the Munster area compared to the overall model area in order to match the lower transmissivity values obtained from the Munster area wells. Boundaries of the lower-conductivity zone were derived from the adjacent geological faults and contacts mapping;
  - The bottom Nepean layer (model layer 5) represents sandstones of relatively high hydraulic conductivity (same as Layer #4 in the Munster area) at the remaining thickness of the unit. The thickness of this layer varies across the model domain; and,
- Precambrian Bedrock (Model Layers 6 and 7) Below the Nepean layers, the Precambrian metasedimentary rocks were discretized into two layers, where the top layer (model layer 6) represents a moderate-low hydraulic conductivity layer, and the bottom layer (model layer 7) represents a 50 m thick zone of low hydraulic conductivity bedrock. The Precambrian surface (illustrated on Figure 11) and Nepean





Formation thickness (illustrated on Figure 12) was interpolated using data from a limited number of boreholes (mainly from the Blacks Corners, Munster, and Manotick areas) and outcrop areas on the outer edges of the model domain.

#### 4.1.8 **Groundwater Flow Boundaries**

Figure 13 illustrates the flow boundaries used in the groundwater flow model. The eastern boundary follows the Rideau River and is specified in model layer 1 as a constant head boundary at an elevation ranging from 80 masl to 85.5 masl (corresponding to the river elevation along the boundary). Similarly, the western model boundary follows the eastern shore of Mississippi Lake (134.5 masl) and the Mississippi River. Along the Mississippi River, the assigned hydraulic head ranged from 122 masl to 134.5 masl. These boundaries were also assigned as constant head within the layers in contact with the water body (layers 1, 2, or 3 depending on the location). The northern and southern boundaries were also defined as constant head boundaries using values obtained from the groundwater elevations in the bedrock aquifer map (Figure 5).

The Jock River flows through the central area of the model in a southwest to northeast direction. A river boundary was assigned to the Jock River based on its elevation and using a conductance value of 200,000  $m^2/d$ .

Four large marshes and wetland areas found within the study area were specified as constant head boundaries. Unique constant head values representing the mean water level elevations (derived from topographic mapping) or non-unique constant head values selected based on topography and on the inferred slope of the Jock River were used to simulate the wetland areas.

The base of the model (the Precambrian rock below model layer 7) is defined as a "no flow" boundary.

#### 4.1.9 **Pumping Wells and Water Takings**

Water demand is primarily a function of population and predicting future demand requires making assumptions about the rate of growth. The average water demand for the Richmond development was assumed to be 1,630 m<sup>3</sup>/d. A detailed calculation of the forecast water demands for the Richmond well system is provided in Section 1.3. For the purposes of the groundwater flow model it was assumed that the forecasted average demand would be supplied evenly between PW08-01 and PW09-01. It should be noted that due to their close proximity (< 3m), wells PW09-01 and PW09-02 were considered as a single well for the purpose of the modelling exercise.

In addition to the Richmond wells, the PTTW database contains other significant water takings within the study area. The Kings Park communal well system is located in the northeastern portion of the Village, approximately 1.5 km northwest of the Richmond wells. This system consists of two wells, RW1 (66 m deep and cased to 19.2 m) and RW2 (61 m deep and cased to 19.5 m). The wells penetrate limestone and dolomite and are completed as open holes in the upper portion of the underlying Nepean Formation sandstone. The permitted capacities of RW1 and RW2 are each 1,310 m<sup>3</sup>/d. Actual pumping rates from these wells are much lower than the maximum permitted rate; historical water use data indicates the average pumping rate is 210 m<sup>3</sup>/d for the Kings Park system. Based on information provided by the City of Ottawa, which projects a zero-percent increase in population for Kings Park, current water demands for the Kings Park system are expected to continue within the time frame considered as a part of the current study.

The Hyde Park communal well system is located approximately one kilometre north of the Richmond wells. This system is comprised of two wells referred to as TW1 (the supply well) and TW2 (the backup well). TW1 and TW2 are completed to depths of 83.8 mbgs and 92 mbgs, respectively, and are completed as open holes



through the Oxford, March, and Nepean Formations. The permitted pumping rate for the system is 576 m<sup>3</sup>/d, though the typical current groundwater usage for this system is approximately 30 m<sup>3</sup>/d based on 2010 usage data. The forecast estimation of average day groundwater usage for the system at full build out is 265 m<sup>3</sup>/d.

The community of Munster, located approximately eight kilometres west-southwest of the Richmond wells, is serviced by two wells (MW1 and MW2). MW1 and MW2 are respectively 116 m and 122 m deep, and are cased to a depth of 29 m below ground surface. The permitted capacities of wells MW1 and MW2 are 980 and 1,181 m<sup>3</sup>/d, respectively. Over the 1998-1999 period, groundwater extraction for this system averaged 422 m<sup>3</sup>/d. The projected increases in pumping rates for these wells based on population growth estimates amount to 5 % above the average rates. For the purposes of this study, the forecast pumping rates used in the 2003 assessment (277 m<sup>3</sup>/d at MW1 and 187 m<sup>3</sup>/d at MW2) were maintained.

A PTTW was identified for a golf course located approximately 14 km west-northwest of the Richmond wells. No usage data was obtained, however, given the location of this well relative to the RV well system, it is anticipated that simulating this usage would not affect the results for the areas of interest of this study. In the absence of actual pumping data, numerical simulations were completed using an assumed pumping rate for the golf course well to confirm that this is the case.

Simulated forecast pumping rates for all wells are summarized in Table 4.

#### 4.1.10 Recharge

Figure 14 illustrates the distribution of recharge used in the groundwater flow model. Three separate recharge zones were defined to reflect the variability in surficial materials in the model area. Clayey overburden material east of the Richmond area was assigned an infiltration rate of 5 mm/y, till and rock outcrop areas west of Richmond were assigned an infiltration rate of 15 mm/y, and a stratified ice contact deposit located northeast of Richmond was assigned an infiltration rate of 200 mm/y. These values were estimated from professional judgement and through the model calibration process. Alternate infiltration configurations were also tested during the parameter variation simulations.

#### 4.1.11 Model Parameterization

Figure 15 illustrates the model hydraulic conductivity distribution, as defined below.

#### **Overburden (Layer 1)**

The overburden material in the model area is comprised primarily of till with some sand and clay areas. There is no detailed differentiation of the overburden materials in areas of the model, and as such the hydraulic parameters are defined globally for the entire overburden layer. Except for bedrock outcropping areas, a unique value of  $5x10^{-7}$  m/s was used in the model. In bedrock outcropping areas, a minimum thickness of 1 m was used and the hydraulic conductivity value of the underlying bedrock layer (based on geological maps) was manually assigned. The effective porosity of the overburden layer was conservatively assigned at 0.25, typical of a sandy silt to silty sand till.

Hydrogeological investigation of the overburden materials local to the Mattamy/RV area was completed previously (Golder, 2010b). Results from hydraulic testing of the silty clay and fine sand overburden materials identified a range in hydraulic conductivity between approximately  $1 \times 10^{-6}$  m/s and  $1 \times 10^{-5}$  m/s. It should be noted that these data were not available at the time of the 2003 assessment, and as such were not used in the original model development. The range in measured hydraulic conductivity for the overburden local to the Richmond





development is higher than the simulated hydraulic conductivity for this unit, though the original simulated value was maintained, as the measured data represent only a small portion of the modelled area.

#### Upper Paleozoic Formations (Layer 2)

The hydraulic conductivity of the dolostones, limestones and shales (Bobcaygeon, Gull River, Rockcliffe, Oxford and March Formations) which overlie the Nepean aquifer range from  $2x10^{-11}$  m/s to  $9x10^{-4}$  m/s based on the results of hydraulic testing of these units. Except for the Nepean bedrock outcropping areas, a unique value of  $5x10^{-7}$  m/s was used in the model. In the Nepean bedrock outcropping areas (based on geological maps), a minimum thickness of 1 m was used and the hydraulic conductivity value of the underlying Nepean bedrock layer was manually assigned. Based on the reported joint spacing and on the bulk hydraulic conductivity used in the model, an effective porosity of 0.001 was estimated and assigned to this layer. These values were varied during the sensitivity analysis.

#### Nepean Formation (Layers 3, 4 and 5)

Based on the results of pumping tests completed on the Richmond wells (PW08-1, and PW09-1), the estimated transmissivity of the Nepean formation at this location is estimated to be between 500 m<sup>2</sup>/d and 800 m<sup>2</sup>/d (Golder, 2011). Assuming that the thickness of the Nepean Formation in this area is approximately 45 m, this corresponds to a range in hydraulic conductivity of  $1 \times 10^{-4}$  m/s to  $2 \times 10^{-4}$  m/s.

Pumping tests completed at wells MW1 and MW2 in Munster indicate bulk bedrock transmissivities in the range of 4 to 370 m<sup>2</sup>/d, with an arithmetic mean of 81 m<sup>2</sup>/d and a geometric mean of 24 m<sup>2</sup>/d. Assuming that the Nepean Formation is the principal contributing aquifer and that its thickness in the Munster area is in the order of 40 metres, this corresponds to a range in hydraulic conductivity of  $1.3 \times 10^{-6}$  m/s to  $1 \times 10^{-4}$  m/s. Also, Packer testing performed in Munster municipal well MW1 indicated a very high hydraulic conductivity (on the order of  $1 \times 10^{-4}$  m/s) in the uppermost portion of this formation.

Pumping tests completed at RW1, RW2 and HP1 (Kings Park/Hyde Park) indicate bulk bedrock transmissivities in the range of 30 to 658 m<sup>2</sup>/d (see Table 2) with an arithmetic mean of 304 m<sup>2</sup>/d and a geometric mean of 216 m<sup>2</sup>/d. Again, assuming that the Nepean Formation is the principal contributing aquifer and that its thickness in the Richmond area is in the order of 45 metres, this corresponds to a range in hydraulic conductivity of  $7.8 \times 10^{-6}$  m/s to  $1.7 \times 10^{-4}$  m/s.

Transmissivity values from pumping tests performed in the Nepean Formation in the Blacks Corners area vary from 148 to 397 m<sup>2</sup>/d with an arithmetic mean of 204 m<sup>2</sup>/d and a geometric mean of 188 m<sup>2</sup>/d. Again, assuming that the Nepean Formation is the principal contributing aquifer and considering a recorded thickness of 50 m in the Blacks Corners area, this corresponds to a range in hydraulic conductivity of  $3.4 \times 10^{-5}$  m/s to  $9.2 \times 10^{-5}$  m/s.

As indicated previously, the Nepean aquifer was discretized into three separate layers in order to better match the site/aquifer specific data and knowledge. Therefore, a hydraulic conductivity value of  $4x10^{-4}$  m/s and a thickness of 1 m were assigned to the Nepean top layer (model layer 3) which is conceptualised as fractured sandstones of relatively high hydraulic conductivity. The Nepean middle layer (mode layer 4) was given a hydraulic conductivity value of  $4x10^{-4}$  m/s and a constant thickness of 4 m through all of the model area except for the Munster area where a value of  $1x10^{-5}$  m/s was assigned in order to match the Munster lower transmissivity values; boundaries of this "lower conductivity" zone are derived from the adjacent geological/faults contacts. The bottom Nepean layer (model layer 5) was assigned a hydraulic conductivity value of  $1x10^{-5}$  m/s and the remaining thickness of the unit. The effective porosity was assumed to be related to the hydraulic





conductivity. Therefore, based on an assumed fracture density of 100/m, any Nepean layer or portion of layer with a hydraulic conductivity value of  $4x10^{-4}$  m/s was given an effective porosity value of 0.017. The rest of the Nepean material ( $1x10^{-5}$  m/s) was considered less fractured and was accordingly given an effective porosity value of 0.001. These values were varied during the sensitivity analysis.

#### Precambrian (Layers 6 and 7)

Information on the hydraulic characteristics of the Precambrian metasedimentary rocks was not available in the study area. Based on values reported in surrounding areas, the upper Precambrian layer (model layer 6) was given a K value of  $1 \times 10^{-7}$  m/s and an effective porosity of 0.001. The lower Precambrian layer (model layer 7), which was considered as less conductive and less fractured, was given a K value of  $1 \times 10^{-8}$  m/s with an effective porosity of 0.0001.

#### Faults

Information on the hydraulic characteristics of the faults in the bedrock was not available in the study area. It is not known whether the faults represent barriers or conduits for groundwater flow. However, and as indicated earlier, available information pertaining to the dissolved contaminant plume in the Beckwith/Blacks Corners area suggests that faults in this sector would be transparent to groundwater migration. The bedrock faults have therefore not been represented as independent hydrostratigraphic units in the groundwater model.

The anisotropy ratio of all hydrostratigraphic units was assumed to be 1:1 (i.e. the vertical hydraulic conductivity is assumed to be equal to the horizontal hydraulic conductivity).

#### 4.1.12 Parameter Variation Simulations

It should be recognized that there is inherently some uncertainty associated with the capture zones forecast by a calibrated groundwater model. These uncertainties stem from limitations in the available subsurface information and can be related to variability in the aquifer properties (e.g., hydraulic conductivity; porosity) or uncertainties with the conceptual model (e.g., groundwater-surface water interactions; location of flow boundaries; recharge rates; continuity in aquitards; direction of regional groundwater flow; simplification of fracture flow systems into bulk hydraulic conductivity (EPM) approaches). To gain some understanding of the potential impact of this uncertainty in the groundwater model forecasts, a sensitivity analysis was completed; the compilation of which when overlaid effectively increases the spatial coverage of each time-of-travel capture zone from those generated using the calibrated model parameters.

The sensitivity analysis was comprised of a series of 21 steady-state groundwater flow simulations, as summarized in Table 5. These simulations considered variability in: the hydraulic conductivity of the upper Paleozoic Formations (model layer 2), the hydraulic conductivity of the Nepean Formation model layers, recharge rates, the effective porosity of the upper Paleozoic Formations, and combinations of these parameter changes.

### 4.1.13 Capture Zone Delineation

As described earlier, the capture zones for the Richmond, King's Park and Munster supply wells were determined using MODPATH by releasing groundwater particles at the pumping wells and backwards tracking them to their source. The time-related capture zones are subsequently derived from this analysis represent a two-dimensional projection of the particle outlines to ground surface. The final capture zones reflect the combined area resulting from the calibrated model simulation and from the various sensitivity runs. Both the Munster and King's Park well systems have two sets of capture zones as these wells are open to both the shallow and deep bedrock aquifers.





## 4.2 **Groundwater Vulnerability Mapping**

Groundwater vulnerability mapping was performed over the area of the model domain using the Intrinsic Susceptibility Index (ISI) method. The ISI method provides a quantitative measure of the degree of protection afforded by the overlying geological material: the higher the index, the greater the degree of protection. The index is calculated for each discrete geological unit by multiplying a "K" factor by the thickness of the layer. A table of K factor values for many geological materials was prepared as a part of the groundwater vulnerability guidance documentation.

As described previously, the Richmond wells are equipped with grouted steel casing to a depth of approximately 45 m and draw water from the Nepean Formation sandstone. Geological materials overlying (and therefore "protecting") this aquifer include the Paleozoic dolomites, limestones, and shales of the March, Oxford, Rockcliffe, Gull River, and Bobcaygeon Formations, and the overlying overburden. Thus, the ISI was evaluated as follows:

$$ISI = 3B_1 + 4B_2$$

where  $B_1$  and  $B_2$  represent the respective thicknesses of the overburden and Paleozoic formations, as interpolated based on the available data (a description of the development of the geological surfaces is included in the 2003 report). The K-Factors of 3 for the overburden and 4 for the bedrock were chosen directly from Table 3.1 in Appendix 3 of the Draft Guidance Module 3; the overburden is conservatively represented as "silty sand", and the Paleozoic formations that overlie the Nepean Formation are represented as "limestone/dolostone".

## 4.3 Intrinsic Vulnerability Scoring

Intrinsic vulnerability scores were calculated based on the results of the groundwater vulnerability mapping and WHPA zone delineation, pursuant to the Technical Rules: Assessment Report Table 2a (ISI). The vulnerability within each of the WHPAs is categorized as high (<30), medium (30 to 80) or low (>80).

## 4.4 Uncertainty Assessment

The zones of uncertainty (i.e. high uncertainty/low confidence, low uncertainty / high confidence) associated with the WHPA delineation and vulnerability scoring qualitatively evaluated in consideration of the quantity and quality of hydrogeological information, the were reasonableness of model calibration compared to available data, and the consistency and repeatability of the parameter variation simulations. Ultimately, zones of uncertainty were delineated based on professional judgement.

## 4.5 Threats Assessment

A threats assessment for the existing King's Park and Munster wells was previously undertaken by Dillon Consulting on behalf of the Mississippi-Rideau Source Protection Region. In accordance with direction from the Source Protection Office of the Ministry of the Environment, only threats that could be considered significant according to the established methodology require evaluation. Significant threats can only occur in two general circumstances:

- 1) In an area where the intrinsic vulnerability score is 8 or 10; or,
- 2) For dense non-aqueous phase liquids (DNAPL's) within WHPA Zones A, B and C.

Potential significant threats to the Mattamy/RV wells were assessed based on these circumstances.



## 5.0 RESULTS

## 5.1 Groundwater Flow Modelling

Calibration of the groundwater flow model was originally completed as a part of the 2003 wellhead protection study. This involved the adjustment of recharge rates to the different overburden and bedrock outcrop zones, the adjustment of the hydraulic conductivities of the various overburden and bedrock units, and the adjustment of the boundary conditions until the simulated groundwater elevations and flow directions compared reasonably well to the observed conditions. The calibration process involved two steps. First, transient simulations were run with adjustments to the three Nepean model layer hydraulic conductivities until a good match was obtained between the observed drawdown measurements obtained during the 2002 pumping testes conducted at the Kings Park and Munster wells. The second step involved adjustments to the recharge rates to the different overburden and bedrock outcrop zones, hydraulic conductivities of the remaining overburden and bedrock units, and the boundary conditions until a reasonable match between the steady state simulated groundwater elevations in the Nepean top layer and the recorded groundwater elevations for bedrock wells completed into the Nepean sandstone.

Calibration of the groundwater flow model involved the adjustment of recharge rates to the different overburden and bedrock outcrop zones, the adjustment of the hydraulic conductivities of the various overburden and bedrock units, and the adjustment of the boundary conditions until the simulated groundwater elevations and flow directions compared reasonably well to the observed conditions. Figure 16 shows the simulated bedrock groundwater elevations following calibration of the model. The simulated groundwater elevations indicate the overall regional groundwater flow direction is towards the east (Rideau River). Local to the Richmond area, groundwater flow is affected by the presence of the Jock River. In general the simulated groundwater flow patterns are consistent with the inferred groundwater elevation map shown on Figure 5.

In addition to reviewing the regional groundwater flow patterns simulated by the model, the static water levels recorded in the MOE WWIS were utilized as discrete points of comparison for the steady-state calibration of the model. A total of 982 calibration points were used following a QA/QC process which removed spurious and/or suspect data from the database (i.e., as defined by wells with a location or elevation accuracy code of 6 or greater, or where reported groundwater elevations were unreasonably high or low compared to nearby data points). It should be noted that MOE WWIS data collected subsequent to the 2003 assessment were included in the calibration process. Figure 17 shows a plot of calibration and provides calibration statistics. Observations made with respect to this figure are summarized below:

- Generally, the simulated groundwater levels compare reasonably well with the measured groundwater levels. Following the "trial-and-error" calibration process, the residual mean was 0.53 m, the absolute residual mean was 3.5 m, and the normalized RMS error was 5.5%;
- There is not a strong bias in simulated groundwater elevations either above or below the historical measured values; and,
- The dense cluster of data points plotted at an observed elevation of 85 masl reflects observations densely clustered in the Manotick area, while the cluster of wells plotted at about 135 masl reflect monitoring locations near the Mississippi Lake and River. This trend is considered to be a result of underlying variability in the topographic data in the MOE WWIS database.



The horizontal hydraulic conductivities that provided the best fit during model calibration (described previously in Section 4.1.11) lie within, or close to, the boundaries of the range of estimates described in Section 3.0, and are considered reasonable estimates of hydraulic conductivity on a regional scale.

A 72-hour pumping test was conducted at PW09-1 between September 27 and September 30, 2011 (Golder, 2011), and the results from this test were assessed with the calibrated groundwater model as a means of independently verifying the model parameterization. The regional model used in this study over-predicts the drawdown that was observed during the 72-hour pumping test at PW09-1. The bulk transmissivity used in the model is more representative of values derived from the tests completed in the Nepean formation in other locations within the study area (summarized previously in Section 3.0). The transmissivity predicted by the test at PW09-1 is higher, and may be more representative of a localized high conductivity zone. Additional testing would be required to verify this interpretation.

As a part of the 2003 assessment, the groundwater flow model calibration was compared against additional pumping tests conducted at the Kings Park and Munster well systems. Further details can be found in the 2003 assessment report.

The calibrated groundwater flow model provides a reasonable understanding of the groundwater flow conditions within the study area. Through the calibration process it was found that the hydraulic conductivities of the geological units are in good agreement with the site-specific information. The calibrated model values therefore represent suitable estimates for the use in developing the theoretical capture zones for the Richmond wells under forecast pumping rates. Table 3 provides a summary of the input parameters and model details for the final calibrated model.

## 5.2 Wellhead Protection Areas and Vulnerability Mapping

As described earlier, the capture zones for the Richmond wells were determined by activating the pumping wells in the calibrated groundwater flow model (given the forecast pumping rates described in Section 1.3) and releasing groundwater particles (using MODPATH) at the pumping wells, which are backward-tracked in the direction of their simulated flow paths. The time-related capture zones that are subsequently derived from this analysis represent a two-dimensional projection of the particle outlines to ground surface. This process was completed for the calibrated groundwater flow model and for each of the parameter variation simulations (for a total of 22 simulations). The projected particle traces from all simulations were combined to form the capture zone areas. The 2-year (WHPA Zone B), 5-year (WHPA Zone C), and 25-year (WHPA Zone D) time of travel capture zones for wells PW-08 and PW-09 as predicted using this method are mapped against the calculated intrinsic vulnerability of the aquifer on Figure 18 (ISI method). Also shown on this figure is the Zone A Pathogen Security / Prohibition Zone (100 m radius around well). Review of the figure allows the following observations:

Due to the proximity of the Jock River, which influences local groundwater flow patterns, the capture zones extend in two directions. The main arm of the capture zones extends towards the west-northwest, in a direction upgradient of regional groundwater flow, through the bedrock, and terminates approximately 14km from the wells in the area of a large wetland located on a topographic high. This arm is approximately 5.5 km in width. The second arm extends approximately 6 km towards the south, beneath the Jock River, and terminates under the large wetland area located south of Richmond. This arm reaches a maximum width of approximately 2 km; and,





Given the thickness of the units overlying the Nepean Aquifer (approximately 40 to 90 m of limestone/dolostone, and 0 m to 15 m of overburden), and the K factors (4 for the bedrock, and 3 for the overburden) the calculated ISI values ranged from 160 to 390. The application of the ISI vulnerability scoring, as described in Section 4.3, therefore resulted in a low groundwater vulnerability score (i.e. ISI>80) over the model domain. ISI values calculated for the Kings Park and Munster lower aquifer as a part of the previous studies were similar to those calculated in the current assessment. The 2008 report notes a range in ISI values of 180 to 350 for these systems, though K factors used for the 2008 assessment were slightly less conservative. Note that this does not apply to the upper aquifer, as the Richmond Village wells do not draw water directly from this aquifer.

Through the parameter variation simulations it was determined that the parameters having the strongest influence on the capture zone splitting (i.e. the "two-arm" configuration) include: the contrast in hydraulic conductivity of the Nepean upper and lower hydrostratigraphic units, variations in hydraulic conductivity of the upper Paleozoic formations, and adjustments to the surface recharge parameter.

## 5.3 Intrinsic Vulnerability Scoring

Figure 19 illustrates the intrinsic vulnerability scoring within the WHPAs for the Richmond wells. As shown on the figure, the calculated vulnerability score ranges from 10 (limited to WHPA Zone A), to 2 in WHPA Zone D.

## 5.4 Uncertainty Analysis

The uncertainty associated with the vulnerability scores within the WHPA was determined in accordance with Draft Guidance Module 3, and is shown on Figure 20. Uncertainty was categorized as either high or low. Based on professional judgement, the areas encompassed within the 2 year WHPA were categorized as "low uncertainty", where the remaining areas within the overall WHPA were categorized as "high uncertainty".

## 5.5 Threats Assessment

The intrinsic vulnerability scores calculated in Section 5.3 and illustrated on Figure 20 are less than 8, with the exception of WHPA Zone A, the 100 metre zone around each well. Therefore significant threats to the Mattamy/RV wells, exclusive of DNAPL's, can only occur within WHPA Zone A. Certain DNAPL's are considered significant threats in WHPA Zones A, B and C with any vulnerability score.

The current land use on the site is agricultural, but this will change to residential/parkland as development proceeds. Based on the current methodology, sewage connections and laterals are defined as wastewater collection facilities and are considered significant threats in areas, such as WHPA Zone A, with an intrinsic vulnerability score of 10. It is our understanding that the MOE has determined that the provision of additional safeguards, such as double lined sewer pipes and more frequent inspections, will be sufficient to mitigate threats from sewage collection infrastructure.

No activities that may propose a significant threat will be permitted within WHPA- Zone A. The draft Source Protection Plan prepared by the MRSPR includes Policies SEW-6-LB and SEW-7-LB-PI-MC. The former includes requirements for the inspection and maintenance of sanitary sewers and related pipes where they are or would be considered a significant drinking water threat. The second policy includes a requirement that approval under the Ontario Water Resources Act includes appropriate measures to manage the threat. This includes minimum construction standards for new or replacement sanitary sewers and related pipes that can be required by the



director of theme as appropriate. As of the date of this report, the Source Protection Office has not confirmed that additional restrictions on land usage within WHPA Zone A are being contemplated.

The storage and handling of DNAPL's is considered a significant threat within WHPA Zones A, B and C (within a 5 year time-of-travel). The capture zones considered as a part of the current study were reviewed and compared to the similar capture zones for the Richmond King's Park model. Both computer and on-ground surveys were used to determine if any potential users of DNAPL's were within the additional 5 year TOT that was not included in the existing assessment report. No additional sources, such as dry cleaners, manufacturing facilities or wood product manufacturers were identified.

## 5.6 Updates to Kings Park and Munster Capture Zones

The Richmond Village wells are completed within the lower aquifer capture zone developed previously for the Kings Park well system. Operation of the Richmond Village wells will draw from groundwater sources that previously would have been extracted by the Kings Park system, and will therefore alter the extents of the Kings Park capture zones. Similarly, the capture zones for the Richmond Village well system presented in Section 5.2 of this report overlap with areas of the capture zone developed previously for the Munster well system, indicating that operation of the Richmond wells may influence the Munster capture areas.

As indicated in Golder's 2009 letter report, both the Kings Park and Munster well systems draw water from the upper and lower aquifers. Because of the potential interaction with the Richmond Village system, upper aquifer and lower aquifer capture zones in for the Kings Park and Munster well systems were updated using the groundwater flow model and methodology described above. Aquifer vulnerability and intrinsic susceptibility indices were again delineated using the updated capture zones.

Results of the reassessment of the Kings Park and Munster well assessments that considered operation of the Richmond Village system are presented in Appendix C, summarized as follows:

- Figure C1 and Figure C2 illustrate the upper aquifer particle traces and capture zones for the Munster and Kings Park well systems respectively. Figure C3 illustrates the combined particle traces for the lower aquifer for both wells systems;
- Figure C4 and Figure C5 illustrate the calculated upper aquifer intrinsic vulnerability (i.e. high, medium or low) for the Munster and Kings Park well systems respectively, and Figure C6 illustrates the Combined Aquifer Intrinsic Vulnerability for both aquifers for both well systems;
- Figure C7 and Figure C8 illustrate the combined intrinsic vulnerability scores for the Munster and Kings Park well systems, respectively; and,
- Figure C9 and Figure C10 illustrate the uncertainty associated with the upper aquifer vulnerability scores within the WHPAs for the Munster and Kings Park well systems respectively. Figure C11 illustrates the uncertainty associated with the lower aquifers for both well systems.

In general, the extents of the updated Munster and Kings Park capture zones illustrated in Figures C1 through C3 are similar to those delineated as a part of previous studies (compare to Figures 1 and 2 from the 2009 report for the upper aquifers, and Figure 2.20 from the 2003 report for the lower aquifers). No change is noted in the capture zone delineation for the Munster upper aquifer (Figure C1). As shown in Figure C3, the updated Munster capture zone for the lower aquifer continues to source water from the north west, though the capture





zone is shifted slightly further towards the south west compared to the original assessment. The updated upper aquifer capture zone for the Kings Park system continues in a southerly direction as per the previous assessment, and is similar in terms of lateral extents. Similarly, the updated Kings Park lower aquifer capture zone follows the same two-arm configuration that was delineated as a part of earlier studies. The northern/western arm is shifted slightly further to the north, and the southern/eastern arm is shifted towards the east and south compared to the previous capture zone.





## 6.0 DATA GAPS

Significant data gaps identified in this study include:

- Characterization of the porosity of the hydrostratigraphic units. Assumed values for effective porosity of the overburden and bedrock units are used in estimating the groundwater velocity (and therefore travel time through a given formation) in the current analysis. Testing (isotope analysis) of groundwater samples from the Nepean Formation from the pumping wells would provide a better estimate of the groundwater travel time through the aquifer and could be used to improve the delineation of the WHPAs. The required testing is not covered under the current project scope;
- Characterization of the bedrock fracture network. Additional data relating to the orientation and extent of the bedrock fracture network identified during well drilling and on regional mapping would provide a better understanding of preferential flow pathways through the bedrock, and improve certainty with respect to the size and orientation of the WHPAs. However, a study of the bedrock fracture network would require significant effort, extending well beyond the scope of the current work;
- The threats assessment included the areas within the new WHPA Zone C, but excluded the areas of overlap between the Kings Park capture zones and the new capture zones. These areas should be re-assessed during the next source protection iteration. Similarly, the areas within the updated Kings Park and Munster capture zones that were not included in the previous capture zones have not been reassessed in terms of threats; and,
- Groundwater users within the model domain were identified based on the available information, including PTTW records, though pumping records were not always available. Incorporation of actual pumping rates into the groundwater model could improve overall delineation of the WHPAs.





## 7.0 LIMITATIONS

This report was prepared for the use of Richmond Village (South) Limited. The report, which specifically includes all tables, figures and appendices, is based on data gathered by Golder Associates Ltd., and information provided to Golder Associates Ltd. by others. The information provided by others has not been independently verified or otherwise examined by Golder Associates Ltd. to determine the accuracy or completeness. Golder Associates Ltd. has relied in good faith on this information and does not accept responsibility for any deficiency, misstatements, or inaccuracies contained in the information as a result of omissions, misinterpretation or fraudulent acts.

The assessment of environmental conditions and possible hazards at this site has been made using the results of physical measurements from a number of locations. The site conditions between testing locations have been inferred based on conditions observed at the testing locations. Actual conditions may deviate from the inferred values.

Hydrogeological investigations and groundwater modelling are dynamic and inexact sciences. They are dynamic in the sense that the state of any hydrological system is changing with time, and in the sense that the science is continually developing new techniques to evaluate these systems. They are inexact in the sense that groundwater systems are complicated beyond human capability to evaluate them comprehensively in detail, and we invariably do not have sufficient data to do so. A groundwater model uses the laws of science and mathematics to draw together the available data into a mathematical or computer-based representation of the essential features of an existing hydrogeological system. While the model itself obviously lacks the detailed reality of the existing hydrogeological system, the behaviour of a valid groundwater model reasonably approximates that of the real system. The validity and accuracy of the model depends on the amount of data available relative to the degree of complexity of the geologic formations, the site geochemistry, the fate and transport of the dissolved compounds, and on the quality and degree of accuracy of the data entered. Therefore, every groundwater model is a simplification of a reality and the model described in this report is not an exception.

The professional groundwater modelling services performed as described in this report were conducted in a manner consistent with that level of care and skill normally exercised by other members of the engineering and science professions currently practising under similar conditions, subject to the quality and quality of available data, the time limits and financial and physical constraints applicable to the services. Unless otherwise specified, the results of previous or simultaneous work provided by sources other than Golder Associates Ltd. and quoted and/or used herein are considered as having been obtained according to recognized and accepted professional rules and practices, and therefore deemed valid. This model provides a predictive scientific tool to evaluate the impacts on a real groundwater system of specified hydrological stresses and/or to compare various scenarios in a decision-making process. However and despite the professional care taken during the construction of the model and in conducting the simulations, its accuracy is bound to the normal uncertainty associated to groundwater modelling and no warranty, express or implied, is made.

Any use which a third party makes of this report, or any reliance on, or decisions to be made based on it, are the responsibilities of such third parties. Golder Associates Ltd. accepts no responsibility for damages, if any, suffered by any third party as a result of decisions made, or actions taken based on this report.





## 8.0 CLOSURE

We trust the information presented in this report meets your requirements. Should you have any questions or concerns, please contact the undersigned.

Yours truly, GOLDER ASSOCIATES LTD FESSION ( SNGINEER 5/2/2013 STEPHEN R. WILS N. F. BISHOP RACTISING MEMBER 100162306 0122 Nicholas Bishop, M.S., P.Eng. Stephen Wilson, P.Geo. ONTAR C ROUNCE OF ONTARIO **Geological Engineer** Senior Hydrogeologist/Associa

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# Table 1Richmond Groundwater VulnerabilityWell Details

Well Name	UTM - Easting (m)	UTM - Northing (m)	Depth to Bedrock (m)	Casing Depth (m)	Total Depth (m)	Formation
PW08-1	433874	5003838	3.4	45.72	137.16	Upper Nepean
PW09-1	433898	5003808	0	45.72	70	Upper Nepean
PW09-2	433897	5003811	4.3	45.72	77.72	Upper Nepean

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# Table 2Richmond Groundwater VulnerabilitySummary of Previous Well Testing

Location	Geological Unit	Minimum Transmissivity (m <sup>2</sup> /day)	Maximum Transmissivity (m <sup>2</sup> /day)	Reference
Village of Richmond	Oxford	9	248	Golder (2006)
Village of Richmond	Oxford	5	100	GeoAnalysis (1991)
Village of Richmond	Oxford, March and Nepean	236	236	Golder (2001), Golder (2008a)
Village of Richmond	Oxford, March and Nepean	112	130	Golder (2004), Golder (2008a)
Village of Richmond	Oxford, March and Nepean	256	658	Graham Berman and Associates (1971)
Village of Richmond	Oxford, March and Nepean	279	642	Jacques Whitford (1991)
Munster	Oxford, March and Nepean	54	370	Jacques Whitford (1990)
Manotick	March and Upper Nepean	600	600	Raven Beck (1996)
Village of Richmond	Nepean	328	700	Golder (2011)
Village of Richmond	Nepean	500	800	Golder (2011)

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# Table 3 Richmond Groundwater Vulnerbality Study Summary of Base Case Modelling Parameterization

Basic Model Construction - MODFLOW Grid Details			
	Number of Cells	158242	
	Number of Layers	7	
	Model Top	variable	(defined by DEM)
1	Model Bottom	variable	(defined by geology)
	Nodal Spacing	5 to 500 m	(variable range)
Hydraulic Properties of Model Hydrostratigraphic Unit	ίs		

	Horizontal Hydraulic Conductivity, K <sub>h</sub>	Vertical Hydraulic Conductivity, K <sub>v</sub>	Effective Porosity
	(m/s)	(m/s)	
Overburden	5x10 <sup>-7</sup>	=K <sub>h</sub>	0.25
Upper Paleozoic Formations	5x10 <sup>-7</sup>	=K <sub>h</sub>	0.001
Upper Nepean	4x10 <sup>-4</sup>	=K <sub>h</sub>	0.017
Middle Nepean (except Munster area)	4x10 <sup>-4</sup>	=K <sub>h</sub>	0.017
Middle Nepean (Munster area)	1x10 <sup>-5</sup>	=K <sub>h</sub>	0.001
Lower Nepean	1x10 <sup>-5</sup>	=K <sub>h</sub>	0.001
Upper Precambrian	1x10 <sup>-7</sup>	=K <sub>h</sub>	0.001
Lower Precambrian	1x10 <sup>-8</sup>	=K <sub>h</sub>	0.0001

#### Model Boundary Conditions

- Variable surficial recharge, ranging from 5 to 200 mm/yr (see Figure 14)

- Refer to Figure 13 for a summary of groundwater model flow boundaries

#### Model Assumptions

- The geological data used in the development of the model was derived from geological information depicted on governmental agency maps, presented in public reports and studies, and filtered from the MOE Water Well Information System.

- Flow is laminar and steady, and is governed by Darcy's Law.

- Regional groundwater flow is simulated using an "equivalent porous media" approach

- Hydraulic heads are vertically averaged within a given model layer.

- A 1:1 horizontal to vertical anisotropy ratio was assumed.

- Except for the middle Nepean layer (Layer #4), a homogeneous hydraulic conductivity was applied for each hydrostratigraphic unit. Spatial variation of material properties within Layer #4 was based on faulted geological contacts reported northeast and southwest of Munster.

- Modelling and capture zone calculations were done at steady-state and therefore the predictions reflect average long-term conditions based on historical data. The potential effects of short-term "out of the ordinary" conditions (floods, severe rainfalls, etc.) were not considered neither the potential effects of long-term climate changes.

- The capture zones that delineate the WHPAs were derived by using a forecast pumping rate based on the estimated build-out of the Richmond Development

- Results assume that no other major water takings occur in the aquifer system that would change the flow directions or water balance near the capture zone of the Kings Park and Munster wells

- There is no differentiation in the overburden units.

- Recharge estimates reflect deeper recharge and discharge characteristics of the groundwater flow system, and do not account for shallow infiltration and discharge to intermittent streams (i.e. interflow).

- Mapped faults and faulted geological contacts were considered to be transparent to groundwater migration and were not explicitly included in the model

- Major rivers and large wetland areas were considered to potentially influence the deep groundwater flow and were included in the model

- A "regionalized" approach to model calibration was employed, such that parameter values were established for the hydrostratigraphic units on a regional scale. Minor, local variations in hydraulic conductivity (which might locally appear to improve the calibration at specific monitoring wells and therefore reduce the overall statistical measure of the calibration error) were not simulated.

- The most recent groundwater elevation data available were used in the calibration process. These are assumed to approximate steadystate conditions.

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# Table 4Richmond Groundwater VulnerabilitySummary of Simulated Well Pumping Rates

Well Name	Simulated Forecast Pumping Rate (m <sup>3</sup> /d)	Formation
Kings Park Well 1	137	March / Nepean
Kings Park Well 2	84	March / Nepean
Munster Well 1	277	March / Nepean
Munster Well 2	287	March / Nepean
PW08-1	815	Upper Nepean
PW09-1	815	Upper Nepean
PW09-2	0	Upper Nepean

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# Table 5 Richmond Groundwater Vulnerbality Study Summary of Modelling Scenarios

	Hydraulic Conductivity (m/s)		Recharge Rates (mm/yr)			Porosity (-)			
Run	Oxford Formation	Upper Nepean and Middle Nepean (Except Munster Area)	Lower Nepean and Middle Nepean (Munster Area)	Clayey Material	Till and Bedrock Outcrops	Stratified Ice Contact Sand and Gravel	Oxford Formation	Upper Nepean and Middle Nepean (except Munster Area)	Lower Nepean and Middle Nepean (Munster Area)
1	5.0E-07	4.0E-04	1.0E-05	5	15	200	0.001	0.017	0.001
2	2.5E-07	4.0E-04	1.0E-05	5	15	200	0.001	0.017	0.001
3	1.0E-07	4.0E-04	1.0E-05	5	15	200	0.001	0.017	0.001
4	1.0E-06	4.0E-04	1.0E-05	5	15	200	0.001	0.017	0.001
5	2.5E-06	4.0E-04	1.0E-05	5	15	200	0.001	0.017	0.001
6	5.0E-07	2.0E-04	1.0E-05	5	15	200	0.001	0.017	0.001
7	5.0E-07	4.0E-04	5.0E-06	5	15	200	0.001	0.017	0.001
8	5.0E-07	4.0E-04	2.0E-05	5	15	200	0.001	0.017	0.001
9	5.0E-07	4.0E-04	1.0E-05	2.5	15	200	0.001	0.017	0.001
10	5.0E-07	4.0E-04	1.0E-05	10	15	200	0.001	0.017	0.001
11	5.0E-07	4.0E-04	1.0E-05	5	7.5	200	0.001	0.017	0.001
12	5.0E-07	4.0E-04	1.0E-05	5	30	200	0.001	0.017	0.001
13	5.0E-07	4.0E-04	1.0E-05	5	15	50	0.001	0.017	0.001
14	2.5E-07	4.0E-04	1.0E-05	2.5	15	200	0.001	0.017	0.001
15	1.0E-06	4.0E-04	1.0E-05	10	15	200	0.001	0.017	0.001
16	2.5E-07	4.0E-04	1.0E-05	5	7.5	200	0.001	0.017	0.001
17	1.0E-06	4.0E-04	1.0E-05	5	30	200	0.001	0.017	0.001
18	5.0E-07	4.0E-04	1.0E-05	5	15	200	0.01	0.017	0.001
19	5.0E-07	4.0E-04	1.0E-05	5	15	200	0.001	0.017	0.017
20	5.0E-07	4.0E-04	1.0E-05	5	15	200	0.001	0.034	0.034
21	5.0E-07	2.0E-04	2.0E-05	5	7.5	200	0.001	0.017	0.001
22	5.0E-07	2.0E-04	2.0E-05	5	7.5	200	0.001	0.034	0.034

Notes:

- Only parameters varied in the sensitivity analysis have been included in the table

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- Shaded values indicate a change from the base case simulation







	_	
380000	0	LEGEND
/	4000	
	502	POPULATED PLACE
LERO		MATTAMY RICHMOND LANDS
3		RAILWAY
	8	ROAD
-	50200(	RICHMOND VILLAGE BOUNDARY
5		SURFICIAL GEOLOGY
24		ORGANIC DEPOSITS
2		SAND DUNES
	6000	FLOODPLAINS, SAND, SILT, CLAY
1	501	FLUVIAL TERRACES, SAND, SILT
X		REWORKED MARINE SEDIMENTS
~		BEACH FORMATIONS
	8	SAND, REWORKED GLACIOFLUVIAL
- 20	50120(	DELTAIC AND ESTUARIAN DEPOSITS
1		MARINE DEPOSITS, CLAY, SILT
13		EROSIONAL TERRACES
		GLACIOFLUVIAL DEPOSITS
102	8000	TILL, PLAIN
ROM	500	TILL, DRUMLINIZED
		TILL, HUMMOCKY TO ROLLING
		PALEOZOIC BEDROCK
ROAD	8	PRECAMBRIAN BEDROCK
$\gamma \setminus -$	50040	WATER
5		
		NOTE
X	00000	THIS FIGURE IS TO BE READ IN CONJUNCTION WITH THE ACCOMPANYING
A.	50(	GOLDER ASSOCIATES LIMITED REPORT NO. 11-1127-0134
		REFERENCE
		1. CANVEC PROVIDED BY HER MAJESTY THE QUEEN IN RIGHT OF CANADA DEPARTMENT OF NATURAL RESOURCES
1º	8	2. BÉLANGER, J. R., URBAN GEOLOGY OF THE NATIONAL CAPITAL AREA, GEOLOGICAL SURVEY OF CANADA, OPEN FILE D3256, 2001
all	49960	PROJECTION: TRANSVERSE MERCATOR DATUM: NAD 83 COORDINATE SYSTEM: UTM ZONE 18
ROI		
0		5 0 5
A		SCALE 1:150,000 KILOMETRES
1	92000	RICHMOND VILLAGE
	496	GROUNDWATER VULNERABILITY
16		TITLE
		SURFICIAL GEOLOGY
C	00	PROJECT No. 11-1127-0134 SCALE AS SHOWN REV. 0.0
3800	49880	Golder GS BJ 15 Feb. 2012 GIGV BJ 15 Feb. 2012 FIGURE 3
33000	<b>~ ~</b> ·	Ottawa, Ontario REVIEW 7/2/16



380000	D	LEGEND
/	4000	POPULATED PLACE
	502	MATTAMY RICHMOND LANDS
LERO		RAILWAY
B		ROAD
-\	8	RICHMOND VILLAGE BOUNDARY
-	50200C	BEDROCK GEOLOGY
5		QUEENSTON FORMATION
	1	CARLSBAD FORMATION
el l	1	BILLINGS FORMATION
	6000	EASTVIEW FORMATION
	501	LINDSAY FORMATION
X	1	VERULAM FORMATION
1	1	BOBCAYGEON FORMATION
	9	GULL RIVER FORMATION
	01200	ROCKLIFFE FORMATION
	4)	OXFORD FORMATION
		MARCH FORMATION
		NEPEAN FORMATION
~	8000	COVEY HILL FORMATION
IOAL	200	GRANITIC (QUARTZ-RICH)
	1	SYENITIC (QUARTZ-POOR)
		BASALTIC ORIGIN, AMPHIBOLITE
OAD	Q	DIORITE, GABBRO
$\langle \cdot \rangle$	00400	GRANITIC ORIGIN, PARAGNEISS
	.,	NON CARBONATE, QUARTZITE
		CARBONATE, MARBLE
		DYKES, PEGMATITE
/.	0000	NOTE
	500	THIS FIGURE IS TO BE READ IN CONJUNCTION WITH THE ACCOMPANYING
-		REFERENCE
$\setminus$		CANVEC PROVIDED BY HER MAJESTY THE QUEEN IN RIGHT     OF CANADA DEPARTMENT OF NATURAL RESOURCES
1	00	2. BÉLANGER, J. R., URBAN GEOLOGY OF THE NATIONAL CAPITAL AREA, GEOLOGICAL SURVEY OF CANADA, OPEN FILE D3256, 2001
	49960	PROJECTION: TRANSVERSE MERCATOR DATUM: NAD 83 COORDINATE SYSTEM: UTM ZONE 18
NROB		
10		5 0 5 
		SCALE 1:150,000 KILOMETRES
>	92000	RICHMOND VILLAGE
	49	GROUNDWATER VULNERABILITY
-		TITLE
		BEDROCK GEOLOGY
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	w6	Ollawa, Onlario REVIEW I Stulsz









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### LEGEND

Overburden (5x10<sup>-7</sup> m/s)
Oxford / March Formation (5x10<sup>-7</sup> m/s)
Upper Nepean (4x10<sup>-4</sup> m/s)
Lower Nepean (1x10<sup>-5</sup> m/s)
Upper Precambrian (1x10<sup>-7</sup> m/s)
Lower Precambrian (1x10<sup>-8</sup> m/s)

NOTE

1. THIS FIGURE IS TO BE READ IN CONJUNCTION WITH THE ACCOMPANYING GOLDER ASSOCIATES LTD. REPORT №. 11-1127-0134 2. THIS FIGURE IS NOT TO SCALE.

RICHMOND VILLAGE WELL SYSTEM - GROUNDWATER VULNERABILITY STUDY							
MODEL LAYERING AND HYDROSTRATIGRAPHIC CONCEPTUALIZATION							
<u> </u>	PROJECT No. 1111270134 FILE No. 1111270134-06.dwg						
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### LEGEND

River Boundary

Constant Head Boundary

Inactive Model Area

Richmond Pumping Wells

#### NOTE

1. THIS FIGURE IS TO BE READ IN CONJUNCTION WITH THE ACCOMPANYING GOLDER ASSOCIATES LTD. REPORT No. 11-1127-0134 2. THIS FIGURE IS NOT TO SCALE.

RICHMOND VILLAGE WELL SYSTEM - GROUNDWATER VULNERABILITY STUDY						
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LEGEND
200 mm/yr Recharge
15 mm/yr Recharge
5 mm/yr Recharge
Inactive Model Area
<ul> <li>Location of Richmond Pumping Wells</li> </ul>

#### NOTE

1. THIS FIGURE IS TO BE READ IN CONJUNCTION WITH THE ACCOMPANYING GOLDER ASSOCIATES LTD. REPORT No. 11-1127-0134 2. THIS FIGURE IS NOT TO SCALE.

PROJECT

RICHMOND VILLAGE WELL SYSTEM -GROUNDWATER VULNERABILITY STUDY

TITLE

### MODEL RECHARGE DISTRIBUTION

-	PROJECT No. 11-1127-0134			FILE No. 1111270134-14 dwg		
	DESIGN	NB	Feb. 2012	SCALE	NTS	REV.
Golder	CADD	BJ	Feb 2012	FIGURE 14		
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Ottawa, Ontario	REVIEW	SU	Bun			



Model Layer 7

410000 415000

## LEGEND Overburden (5x10-7 m/s) Oxford / March Formation (5x10<sup>-7</sup> m/s) Upper Nepean (4x10<sup>-4</sup> m/s) Lower Nepean (1x10<sup>-5</sup> m/s) Upper Precambrian (1x10<sup>-7</sup> m/s) Lower Precambrian (1x10<sup>-8</sup> m/s) Inactive Model Area • Location of Richmond Pumping Wells

#### NOTE

1. THIS FIGURE IS TO BE READ IN CONJUNCTION WITH THE ACCOMPANYING GOLDER ASSOCIATES LTD. REPORT No. 11-1127-0134 2. THIS FIGURE IS NOT TO SCALE.

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	Ottowa Ontario	DEVIEW	al.	4171			













## **APPENDIX A**

**Peer Review Correspondence** 





June 11, 2012

Richmond Village (South) Limited 3894 Prince of Wales Drive Ottawa, Ontario K2C 3H2

Attention: Mr. Frank Cairo

#### Peer Review of Groundwater Vulnerability Study Richmond Village Well System, Richmond, Ontario

Dear Mr. Cairo:

This letter summarizes our peer review of the report, titled *Groundwater Vulnerability Study, Richmond Village Well System, Richmond, Ontario*, prepared by Golder Associates Ltd., dated March 2012. The peer review objectives are to:

- Compare the methodology and results of the study with the expectations of the final Technical Rules: Assessment Report, *Clean Water Act, 2006* dated November 2009.
- Assess the approach used to determine the groundwater vulnerability.
- Assess from a scientific viewpoint the results of the groundwater vulnerability study.

#### BACKGROUND

...continued

A new well field has been established for new development on the west side of Richmond, Ontario on property owned by Richmond Village (South) Limited and Richmond Village (North) Limited (together referred to as RV). The new well field will also service development completed by Mattamy Homes Limited (Mattamy). Currently, existing homes in the village are serviced by individual private wells and septic systems. Starting in 2004, studies were completed including a Class Environmental Assessment and subsequent Final Design of a new municipal water supply system, based on communal wells. In 2009, Golder was retained by the Mississippi-Rideau Source Protection Region (MRSPR) to complete a Groundwater Vulnerability Study for the communal water supply system.

130 Dufferin Avenue London, Ontario Canada N6A 5R2 Mail: Box 426 London, Ontario Canada N6A 4W7 Telephone (519) 438-6192 Fax (519) 672-8209

Dillon Consulting Limited



Richmond Village (South) Limited Page 2 June 11, 2012

There are three proposed municipal wells designated PW08-1, PW09-1 and PW09-2. PW08-1 is 137.2 m deep, and has a steel casing extending to a depth of 45.7 m. PW09-1 is 70.0 m deep with a steel casing also extending to a depth of 45.7 m. PW09-2 was installed less than 5 m from PW09-1 with a similar completion depth. The average water demand for the well system was calculated to be 1,630 m<sup>3</sup>/day for the combined RV/Mattamy development, which includes 2,000 single homes (835 L/day/unit) and 350 town homes (720 L/day/unit).

The primary aquifer for the production wells is the Nepean Formation, a regional extensive bedrock formation consisting of alternating beds of sandstone and quartz. Overlying the Nepean Formation in the Richmond area are the March and Oxford Formations. These bedrock units are also aquifers but are less productive than the Nepean Formation aquifer. Generally, two distinct aquifers are present within the area: an upper aquifer, typically defined as the upper 35 m of the Oxford Formation and a lower aquifer consisting of the lower portion of the March Formation and the upper portion of the Nepean Formation. The majority of private residential wells are completed within the upper aquifer and communal wells are installed in the more productive lower aquifer.

In addition to the new production wells for the RV/Mattamy development lands, there are several other large production wells in the Richmond area. These include the Kings Park communal well system, located about 1.5 km north of the RV wells, and the Hyde Park communal well system located approximately 1 km north of the RV wells. The Munster communal wells are located further away from the RV wells, about 8 km west-southwest from Richmond. Also included in the study area is a golf course well located about 14 km west-northwest of Richmond.

#### **GROUNDWATER FLOW MODEL**

The groundwater flow model is based on a conceptual flow model consisting of the overburden and bedrock formations of the Richmond area. The conceptual flow model is based on previous assessments in the area, notably a 2003 groundwater vulnerability study completed for the Kings Park and Munster well systems (Golder, 2003) and a 2010 study that addressed the Hyde Park well system.

A numerical model called MODFLOW was used to complete the groundwater flow modeling. MODFLOW is a widely accepted program and is recognized by the MOE as being a preferred method for delineating wellhead protection zones.



Richmond Village (South) Limited Page 3 June 11, 2012

#### **Boundary Conditions**

Boundary conditions define input/output zones for the model and MODFLOW allows a number of different types of boundary conditions. Report Figure 13 shows the boundary conditions used in the model. The eastern boundary consists of a 'constant head boundary' representing the Rideau River. Similarly, the western boundary is a constant head boundary representing Mississippi Lake and the Mississippi River. The values for the constant head cells varied to represent the actual stage of both the Rideau River and the Mississippi River.

The northern and southern boundaries were also defined as constant head boundaries and assigned constant head values based on groundwater elevations in the bedrock aquifer map.

In addition to the perimeter boundary conditions, other features were represented in the model. These include the use of a 'river boundary condition' to represent the Jock River which flows through the central area of the model and four large marches/wetland areas which were modeled as constant head boundaries.

#### Model Calibration - Hydraulic Conductivity and Recharge

Six hydraulic conductivity zones (report Figure 6 and report Figure 15) were specified in the model: overburden, Oxford / March Formation, Upper Nepean, Lower Nepean, Upper Precambrian and Lower Precambrian.

Recharge was defined in the model based on overburden type and thickness. Three recharge zones were designated (report Figure 14) representing clayey overburden located mainly east of Richmond (5 mm/year), till and rock outcrop areas west of Richmond (15 mm/year) and a stratified ice contact deposit located northeast of Richmond (200 mm/year).

The flow rate assigned to the production wells was  $1,630 \text{ m}^3/\text{day}$ , equally proportioned between PW08-1 and PW09-1 (since PW09-2 is located less than 5 m from PW09-1 it was not distinctly included in the model). A flow rate of 265 m<sup>3</sup>/day, representing full build out, was used for the Hyde Park well system and a 210 m<sup>3</sup>/day was used for the Kings Park well system.



Richmond Village (South) Limited Page 4 June 11, 2012

The primary calibration procedure used for the model development was to complete a steady-state calibration using 982 calibration wells from the MOE Water Well Information System (WWIS). The calibration plot shown on report Figure 17 indicates a good fit for calibration wells with a normalized root mean square of 5.5% which is considered to be a 'good' result given the variability in the Water Well Record data. It is unclear in the report if the ratio of hydraulic conductivities of the units were kept constant in the calibration process. Likewise, information on how the recharge rates were varied in the calibration process is also not provided.

There is good documentation in the report comparing the results of the calibrated hydraulic conductivity values with measured hydraulic conductivity values determined using other means such as pumping test analyses. Overall, the model construction, boundary conditions and calibration procedure was adequately described in the report.

To define capture zones based on the MODFLOW groundwater flow model, an adjunct computer program called MODPATH was used. A critical parameter used by MODPATH is the effective porosity. Effective porosity assumptions used in the model were: 0.25 for overburden, 0.001 for 0.017 for the upper Nepean and 0.001 for the lower Nepean, and the Oxford /March formations and upper Precambrian layers and 0.0001 for the lower Precambrian layer. Overall, we agree with the approach taken in assessing effective porosity.

Once the groundwater flow model was calibrated, a sensitivity assessment was completed by varying hydraulic conductivity and recharge values. A total of 21 steady-state sensitivity simulations were completed. The defined capture areas are an overlay of all of the sensitivity runs. The sensitivity assessment approach was very comprehensive.

#### **COMMENTS ON THE MODEL**

The following are remarks made on the results of the model.

• A full build-out pumping rate was used for the RV wells. In reality it will be many years until that flow rate is achieved and as such, is considered to be an appropriate conservative assumption. There is some ambiguity in the documentation on the exact pumping values used for all communal wells and a table summarizing the flow rates used in the simulations would have been beneficial.



Richmond Village (South) Limited Page 5 June 11, 2012

- As stated previously, the defined capture zones are the combination of a number of sensitivity simulations of assumed conditions. This approach is considered the best way to assess uncertainty with the model and is considered to be an appropriately conservative approach to establishing the capture zones. It would be interesting to have an illustration of the capture zones based only on the calibrated model. A comparison could then be made between the calibrated capture zones and the ultimate combined capture zones.
- More information is required on how the calibration process was completed (i.e., how the hydraulic conductivity and recharge values were varied in the calibration process).
- Data gaps and limitations are documented in the report. We concur with the data gaps assessment and limitations that are inherent in groundwater modelling.

The concerns expressed in the first three comments above all indicate that the capture zones depicted in report are 'conservatively' large. While such conservativeness is appropriate given the uncertainty associated with groundwater modeling and the lack of good hydrogeological data in areas distant from the wellfield, it is reinforced that the model (and thus capture zone delineation) can be improved with additional study and reduction of data gaps.

#### AQUIFER VULNERABILITY AND VULNERABILITY SCORING

Aquifer vulnerability mapping was undertaken using the Intrinsic Susceptibility Index (ISI) approach. The results from this analysis were combined with the modeled vulnerability zones (WHPA-A, B, C, D) to generate vulnerably scores. The method used is consistent with the requirements of the MOE *Clean Water Act* Technical Rules – Assessment Report. Overall, the results of the assessment appear reasonable; however, Dillon does make the following recommendations:

• Section 4.2/5.2. It is not clear if ISI values were calculated at each MOE water well location within the vulnerability zones (WHPAs), or whether the assessment was performed using interpolated geological/model layer thicknesses. It is recommended that the report further describe how the ISI equation was applied in this study. In addition, it is recommended that the text refer to the actual calculated ranges of ISI values, rather than just identifying that the aquifer had a score >80. A map showing the contoured ISI values would be useful.

....continued



Richmond Village (South) Limited Page 6 June 11, 2012

- Section 4.2/5.2. It is recommended that the report provide further discussion on the significance of the deep casings in these wells as it relates to calculation of the aquifer vulnerability. We note that areas around other municipal bedrock wells (where more shallow casings are used) in other portions of eastern Ontario have been assigned higher aquifer vulnerabilities. It would be useful to reference specific field observations in the local area that support the conclusion that the Nepean Formation is inherently confined and therefore deserving of a low vulnerability ISI value. For example, specific information (if available) on flowing well conditions, large vertical gradients or lack of response in shallow aquifers/surface water features from pumping of the Nepean aquifer would be useful.
- Section 5.3 It would be useful to compare the vulnerability score results to those of the Kings Park system. Since the geology is similar, we would expect that the vulnerability scoring to be similar.

#### **UNCERTAINTY ANALYSIS**

The study concludes that the uncertainty is considered low within Zone A and B, and high within Zone C and D. Dillon is in general agreement with this conclusion; however, it is recommended that the authors expand on this section to provide a rational for this decision. Section 4.4 outlines the key variables that are considered in the uncertainty analysis, however, the report does not comment on how these variables were applied/evaluated in this situation. Further explanation would be beneficial.



Richmond Village (South) Limited Page 7 June 11, 2012

#### CLOSURE

We appreciate the opportunity to assist you with this project. If you have any questions or require clarification regarding our review, please call the undersigned.

Yours sincerely,

#### DILLON CONSULTING LIMITED

Rob Kell, P.Eng., P.Geo. for Darin Burr, P.Geo. Project Manager

RFK:amb

Our File: 12-6360

Document No. 11-1127-0134



December 21, 2010

Darrin Burr, P.Geo. Dillon Consulting 130 Dufferin Avenue London, Ontario N6A 5R2

#### PEER REVIEW OF GROUNDWATER VULNERABILITY STUDY RICHMOND VILLAGE WELL SYSTEM, RICHMOND, ONTARIO

Dear Mr. Burr:

Dillon Consulting (Dillon) was retained to provide a professional peer review of the Golder Associates Ltd. (Golder) report "Groundwater Vulnerability Study, Richmond Village Well System, Richmond, Ontario" (March 2012). The reviewer provided a number of comments in a letter to Frank Cairo dated June 11, 2012, and Golder has prepared the following letter in response to these comments.

The text of each Dillon comment that required a response is presented below, immediately followed by the Golder response. The comments are presented in the same order as in the June letter.

**Dillon** – A full build-out pumping rate was used for the RV wells. In reality it will be many years until that flow rate is achieved and as such, is considered to be an appropriate conservative assumption. There is some ambiguity in the documentation on the exact pumping values used for all communal wells and a table summarizing the flow rates used in the simulations would have been beneficial.

**Response** – The pumping rates used in the forecast model for all groundwater wells are summarized in the table below. This table will be added to the final version of the report.

Well Name	Simulated Forecast Pumping Rate (m <sup>3</sup> /d)	Formation	
Kings Park Well 1	137	March / Nepean	
Kings Park Well 2	84	March / Nepean	
Munster Well 1	277	March / Nepean	
Munster Well 2	287	March / Nepean	
PW08-1	815	Upper Nepean	
PW09-1	815	Upper Nepean	
PW09-2	0	Upper Nepean	

**Dillon** – As stated previously, the defined capture zones are the combination of a number of sensitivity simulations of assumed conditions. This approach is considered the best way to assess uncertainty with the model and is considered to be an appropriately conservative approach to establishing the capture zones. It would be interesting to have an illustration of the capture zones based only on the calibrated model. A comparison could then be made between the calibrated capture zones and the ultimate combined capture zones.

**Response** – A figure is attached showing the base case (i.e. calibrated model) particle pathways overlain with the overall capture zones for the Richmond Village wells. As shown in the figure, the base case capture zones generally occupy the central portions of the overall capture zones, with the exception of the southern extension.

**Dillon** – More information is required on how the calibration process was completed (i.e., how the hydraulic conductivity and recharge values were varied in the calibration process).

**Response** – Model calibration was completed at the time of the original assessment for the Kings Park and Munster wells. The text below, taken from the 2003 wellhead protection study report, will be added to Section 5.1 of the current report to further expand on the calibration process.

Calibration of the groundwater flow model was originally completed as a part of the 2003 wellhead protection study. This involved the adjustment of recharge rates to the different overburden and bedrock outcrop zones, the adjustment of the hydraulic conductivities of the various overburden and bedrock units, and the adjustment of the boundary conditions until the simulated groundwater elevations and flow directions compared reasonably well to the observed conditions. The calibration process involved two steps. First, transient simulations were run with adjustments to the three Nepean model layer hydraulic conductivities until a good match was obtained between the observed drawdown measurements obtained during the 2002 pumping testes conducted at the Kings Park and Munster wells. The second step involved adjustments to the recharge rates to the different overburden and bedrock outcrop zones, hydraulic conductivities of the remaining overburden and bedrock units, and the boundary conditions until a reasonable match between the steady state simulated groundwater elevations in the Nepean top layer and the recorded groundwater elevations for bedrock wells completed into the Nepean sandstone.

**Dillon** – Section 4.2/5.2. It is not clear if ISI values were calculated at each MOE water well location within the vulnerability zones (WHPAs), or whether the assessment was performed using interpolated geological/model layer thicknesses. It is recommended that the report further describe how the ISI equation was applied in this study. In addition, it is recommended that the text refer to the actual calculated ranges of ISI values, rather than just identifying that the aquifer had a score <80. A map showing the contoured ISI values would be useful.

**Response** – The ISI values were calculated using the interpolated geological layer thicknesses, which were developed based on MOE water well records and 'golden spike' borehole data (a description of the development of geological surfaces is included in the 2003 report). Given the thickness of the overlying units above the Nepean Aquifer (approximately 40 to 90 m of limestone/dolostone, and 0 m to 15 m of overburden), the calculated ISI values ranged from 160 to 390.

ISI values calculated for the Kings Park and Munster lower aquifer as a part of the previous studies were similar to those calculated in the current assessment. The 2008 report notes a range in ISI values of 180 to 350 for these systems, though K factors used for the 2008 assessment were slightly less conservative. Note that this does not apply to the upper aquifer, as the Richmond Village wells are equipped with sufficient casing to cut off water from this source.



Text to this effect will be added to the report in Sections 4.2 and 5.2.

**Dillon** – Section 4.2/5.2. It is recommended that the report provide further discussion on the significance of the deep casings in these wells as it relates to calculation of the aquifer vulnerability. We note that areas around other municipal bedrock wells (where more shallow casings are used) in other portions of eastern Ontario have been assigned higher aquifer vulnerabilities. It would be useful to reference specific field observations in the local area that support the conclusion that the Nepean Formation is inherently confined and therefore deserving of a low vulnerability ISI value. For example, specific information (if available) on flowing well conditions, large vertical gradients or lack of response in shallow aquifers/surface water features from pumping of Nepean aquifer would be useful.

**Response** – The reviewer is correct to observe that the aquifer vulnerabilities calculated for other municipal bedrock wells in Eastern Ontario were greater than that assessed for the new Richmond Village wells. It is precisely the reason that additional casing was installed in the RV wells. For example, the municipal wells in Kemptville (North Grenville) and Merrickville were assigned two sets of capture zones and groundwater vulnerabilities, one for the upper aquifer contained within the Oxford Formation and another for the deeper aquifer in the March and Nepean Formations. These municipal wells were constructed with relatively short well casings and were open to the upper aquifer. While the bulk of the water extracted from these wells comes from the deeper, highly transmissive sandstone aquifer, some water originates in the upper aquifer. The ISI values calculated for the upper aquifer are much greater (lower ISI value) due to the lower thickness of material overlying the aquifer. The final aquifer vulnerability maps included both aquifers, with the higher vulnerability of the two shown on the map.

Both the Merrickville and Kemptville wells have since been reconstructed to ensure that only the deep aquifer is being utilized, and the ISI vulnerability scores now reflect the deep aquifer only. The King's Park (Richmond) wells are similar in that both the upper and lower aquifers are connected to the well bore. We understand that the City of Ottawa endeavored to reconstruct these wells to reduce the aquifer vulnerability, but the relatively narrow casing diameter (203 mm) made this effort impractical. The Richmond West wells are equipped with approximately 45 metres of well casing which prevents water from the upper Oxford aquifer entering the well bore.

The heads in the Nepean aquifer within Richmond are above ground surface, and wells constructed within the Nepean are flowing artesian wells. The head in the test wells was observed to be as much as 5 metres above ground surface. Wells completed in the upper aquifer do not flow to the same extent, if at all, and the heads tend to be two to six metres below that in the lower aquifer. High volume pumping of the deeper aquifer can produce minor drawdown in nearby wells completed in the upper aquifer, indicating that there is some hydraulic connection between the units. However, the vertical gradient between the upper and lower aquifer is maintained during higher volume pumping, thus the lower Nepean aquifer is considered to be confined.

**Dillon** – Section 5.3 – It would be useful to compare the vulnerability score results to those of the Kings Park system. Since the geology is similar, we would expect that the vulnerability scoring to be similar.

**Response** – As per the comment above regarding ISI values for the Kings Park and Munster lower aquifer, vulnerability scores for these capture zones were similar to those of the Richmond Village system. The capture zones for the new Richmond Village wells have for the most part displaced the former deep capture zones for the King's Park wells. As mentioned in a previous response, the ISI scores applied in 2012 are somewhat more conservative than applied previously, but would be identical to the King's Park deep scores, as it is the same aquifer being assessed.



As a response to a comment received from the Rideau Valley Conservation Authority (RVCA), the King's Park and Munster capture zones have been reassessed and are included in the amended report.

GE -0 Yours truly, GOLDER ASSOCIATES LTD PRACTISING MER 0:22 Stephen Wilson, P.Geo. ONTAN Associate, Senior Hydrogeologist

#### SRW/NFB/BTB/sg

n:\active\2011\1127 - geosciences\11-1127-0134 richmond village groundwater vulnerability\reporting\letter\_responses to dillon comments 21dec2012.docx

cc: Frank Cairo, Richmond Village (South) Limited

Attachments: Figure: Wellhead Protection Areas and Base Case Particles (Richmond Village Wells)



DILLON

January 9, 2013

Richmond Village (South) Limited 3894 Prince of Wales Drive Ottawa, Ontario K2C 3H2

Attention: Mr. Frank Cairo

#### Peer Review of Groundwater Vulnerability Study Richmond Village Well System, Richmond, Ontario

Dear Mr. Cairo:

Dillon Consulting Limited (Dillon) has reviewed the Golder Associates Ltd. (Golder) December 21, 2012, letter that provided responses to our peer review of the Golder report titled *Groundwater Vulnerability Study, Richmond Village Well System, Richmond, Ontario, March 2012.* Dillon's peer review comments were issued in a letter to Richmond Village (South) Limited dated December 21, 2012.

After review of the Golder responses, and through additional telephone and e-mail communications with Stephen Wilson, P.Geo., of Golder, it is our opinion that Dillon's peer review comments have been addressed.

We appreciate the opportunity to assist you with this project. If you have any questions or require clarification regarding our review, please call the undersigned.

Yours sincerely,

#### DILLON CONSULTING LIMITED

Darin Burr, P.Geo. Project Manager

DTB:amb

Our File: 12-6360

130 Dufferin Avenue London, Ontario Canada N6A 5R2 Mail: Box 426 London, Ontario Canada N6A 4W7 Telephone (519) 438-6192 Fax (519) 672-8209 From: Wilson, Stephen (Ottawa) Sent: January-08-13 10:50 AM To: dburr@dillon.ca Subject: Richmond Village Groundwater Vulnerability Study

Darin,

As requested, the following has been prepared to explain the uncertainty assessment used in the document.

Evaluating Uncertainty in the Assessment.

Assigning an uncertainty rating for specific areas within the capture zones is a somewhat subjective exercise. The assignment of high or low uncertainty within the Wellhead Protection Areas in the Richmond Village model was undertaken with consideration of Appendix 6 of the Assessment Report: Draft Guidance Module 3- Groundwater Vulnerability Analysis (MOE, October 2006).

The decision to assign the areas of low uncertainty to that contained within the two-year TOT capture zone was based on the following:

- High quality data in the form of pumping test information was available near the wells. The data included monitoring well response in both the upper and lower aquifers. This information provided relatively high confidence that the conceptual model correctly represented the aquifer flow systems in the area, as the model was calibrated to the data;
- The modelling approach considered a number of potential scenarios, all of which were considered reasonable. Based on MOE guidance (Appendix 6c – Assigning an Uncertainty Rating for Each Sensitivity Area), it could be argued that the entire capture zone (WHPA-A, B, C and D) could be assigned low uncertainty. However, considering the relatively low quality and density of data beyond the two year time of travel (WHPA-B) as compared to that near the wells, the WHPA-C and WHPA-D zones were assigned high uncertainty. This was a decision made using professional judgement, as per Appendix 6C in the guidance document.

We trust this is sufficient for your needs. If we can be of any assistance please contact the undersigned.

#### Steve

 Stephen Wilson, P.Geo.
 Senior Hydrogeologist/Associate
 Golder Associates Ltd.

 32 Steacie Drive, Kanata, Ontario, Canada K2K 2A9
 T: +1 (613) 592 9600 | D: +1 (613) 287-3286 Ext 3279 | F: +1 (613) 592 9601 | E:
 srwilson@golder.com | www.golder.com

#### Work Safe, Home Safe

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# **APPENDIX B**

**Borehole and Water Well Records** 




First Name Last Name Organization Mailing Address (Street Number/Name) 123 Huntmar Orive Well Eccation Address of Well Location (Street Number/Name) County/District/Municipality County/District/Municipality UTM Coordinates Zone Easting Northing NAD 813	City/Town/Village Rich-model City/Town/Village Rich-mond Municipal Plan and Subjet Number	Address Postal Code K 2 S   B 9 6 7	Well Constructed by Well Owner one No. (inc. area code) 3 & 2 3 4 7 7 secton 3. Postal Code	
General Colour Most Common Material General Colour Most Common Material Brown. Clay. black Shake. white Sandstone black Shake	Other Materials	General Descriptions	Depth (m/t) From To 0 14 1/4 <sup>1</sup> 180 180 195 180 195 185 255 1	and the second sec
Annular Space Depth Set al (m/R) From To To (Malerial and Type) O 150° High Yearly Ce	Volume Placed (m?//?) Ment 103	Results of Well Yield Tee           (well yield, water was:         Draw Do           and sand free         Time Wate           specify         Static           Level         1           ke set at (m/t)         2	ting: wn Recovery rt.evel Time Water Level (min) Mater Level (min) 1 1 2	14 1 1 1
Mithod fof Construction         Øfable Tool       Diamond         Mathod for Constructional       Dating         Rotary (Conventional)       Dating         Boring       Digging         Air parcussion       Digging         Other, specify       Construction Record - Casing         Bunde       Galvarized, Fibreglas, Trickness         Dameter       Construction Record - Casing         Bunde       Construction Record - Fibreglas, Trickness         Offset       From         Offset       Stace	Woll Use     Pumping i       Commercial     Doutation of Devalering       Or Test Hole     Monitoring       Cooling & Air Conditioning     If flowing a final water Supply       To     Status of Well       To     Test Hole       Cooling & Air Conditioning     If flowing a final water Supply       To     Status of Well       To     Test Hole       To     Test Hole       Cooling & Recharge Well     Recomment Well       Downatering Well     Downatering Well	2         3           4         4           4         5           1         5           1         10           1         10           1         10           1         10           1         15           20         20           0         25           nded pump rate         30	2 3 4 5 10 15 20 25 30	
Construction Record - Screen Outside Damier (Plastic, Galvanized, Steel) Stot No From Water Deput	Observation and/or Monitoring Hole     Observation Alteration (Construction) Abandoned, Insufficient Supply Abandoned, other, specify     Other, specify     Hole Danmeter	Interior (Emin / GPM)	40 50 60 p//	
Water found at Depth Kind of Water: Fresh Inteste 73. (m/l) Gas Other, specify Water found at Depth Kind of Water: Fresh Unitests (10 (m/l) Gas Other, specify Water found at Depth Kind of Water: Fresh Unitests 210 (m/l) Gas Other, specify Water found at Depth Kind of Water: Fresh Unitests 210 million Gas Other, specify Business Name of Well Contractor and Well Tachnic Business Address (Street Yumber/Name) 23 With Cheen pd.	an Hifermalion Action and Action	Franktown R	2	- 13
Province Postal Code Business E-mail A Bus Telephone No. (inc. area code) Name of Weil Technician Bus Telephone No. (inc. area code) Name of Weil Technician Bus Telephone No. (inc. area code) Name of Weil Technician Weil Technicaler's Licence No. Significant of Technician and/or Topological Statement No. Significant of Technician and/or Business E-mail A	Meress Last Name, First Name) Chart Name, First Name) Contrictor Date Submitted Contrictor Date Submitted D   0   1 0 M   M   D   0 Ministry's Copy	r's Date Package Delivered Y  Y  Y  V  M  M D   D Date Work Completed Date () () () () () () () () () () () () ()	Alfrituity Use Only Z 103267	

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Measurements recorded in: Matric Mimperial	A 089809	Regulatio	m 903 Ontario Water Resources Ac
Well Owner's Information		PW02	S - Page 1 of 1.
Last Name / Organization Mattamy (	Jack River Limite	E-mail Address	Weli Constructed by Well Owner
123 Huntoner Drive	Ottawa.	Province Postal Code	B9.6138253479
Address of Well Location (Sireet Number/Name)	Township	Loi	Concession
County/District/Municipality Goulbourn	city/Town/Village Richmon	d.	Provinco Postal Code Ontario
NAD 8 3 7 5 50 496 511	Municipal Plan and Sub	at Number	Other
Overburden and Bedrock Materials/Abandonment Seal General Colour Most Common Material	ing Record (see instructions on the Other Materials	e back of this form) General Description	Depth (m
Brown Clay		Packed.	
plackShale		soft	150 150
Slack Shake			195 255
black. granite			755 395 395 450
		1	in the second
Annular Space	Advantation respectively	Results of We	all Yield Testing
From To (Material and Type) O 150° High Parly Rema	(mm) (mm) 103	Clear and sand free	Tinte Water Level Time Water Level (min) (m/R) (min) (m/R)
Jucinfreem		If pumping discontinued, give reason;	Level
		Pump intake set at (m/lt)	2 2
Method of Construction	Well Use	Pumping rate (Imin / GPM)	3, 3
Rotary (Conventional)JettingDomostic Rotary (Reverse)DrivingLivestock	Municipal         Devratering           Test Hole         Monitoring	Duration of pumping hrs + min	5 5
Air parcussion I industrial Clinit, specify I oliver, specify	) Cooling & All Conditioning	Final water lever end of pumping (mill)	10 10
Construction Record - Casing	Status of Well	It flowing give rate (Itrain (GPM))	
Inside Open Hole OR Material Wol Depth (r. Farneter (Galvanized, Fibreglass, Thickness, (onliny Concrete Plastic, Steel) (on (r.)) From	To Water Supply	Recommended pump depth (m/h)	20 20 20 20 20
21/8" Steel 1.88" O	150 <sup>1</sup> Recharge Well	Recommended pump rate (Kmin / GPM)	30 30
	Observation and/or     Monitoring Hole	Well production (Umin / GFM)	40 40
	Construction (Construction)	Elistration? Vyes No	50 50 60 60
Construction Record - Screen	losufficient Supply	Map of We	Il Location
muchar Material Depth (a Depth (a Depth (a Calvanized, Steel) Slot No. From	To Vater Quality To Abandoned, other, specify	610 FF	instructions on the back,
	Other, specify	$  \uparrow \uparrow$	
Water Details       tor found at Depth Kind of Water:     Fresh     Untested       3     (n(0))     Gas     Olner, specify       3er found at Depth Kind of Water:     Fresh     Untested	Hole Diameter Depth (m/tt) Diameter From To (cm/c))	1 Km	n
<u>A (m/h)</u> Gos Other, specify er found at Depth Kind of Water: Frash Untested	255 450 61/8		-
Well Contractor and Well Technician h	nformation	V	



# **APPENDIX C**

**Updates to Kings Park and Munster Capture Zones** 





TAW 🕂	ER SUPPLY WELL
—— 2 Ye	ar Particle Trace
—— 5 Ye	ar Particle Trace
—— 25 Y	ear Particle Trace
Roa	dway
	way
•—— Utilit	y Line
Wate	ercourse, Permanent
— — Wat	ercourse, Intermittent
Build	ding
Wat	er Area, Permanent
Wet	and, Permanent
Mun	ster Hamlet Boundary
Upp	er Aquifer Wellhead Protection Area

### NOTE:

This figure is to be read in conjunction with the accompanying Golder Associates Ltd. report No. 11-1127-0134

### REFERENCE

Base data provided by DMTI Spatial Inc, 2006, NRVIS, 2004, used under license. Projection: Transverse Mercator Datum: NAD 83 Coordinate System: UTM Zone 18



PROJECT

TITLE

### RICHMOND VILLAGE GROUNDWATER VULNERABILITY

### MUNSTER HAMLET UPPER AQUIFER PARTICLE TRACES



PROJECT No. 11-1127-0134		1127-0134	SCALE AS SHOWN	REV. 2		
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GIS	ABD	Jan. 2013				
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REVIEW						



WATER SUPPLY WELL
2 Year Particle Trace
5 Year Particle Trace
25 Year Particle Trace
Railway
Roadway
Utility Line
Watercourse, Permanent
— — Watercourse, Intermittent
Building
Water Area, Permanent
Wetland, Permanent
Village of Richmond Boundary
Upper Aquifer Wellhead Protection Area

### NOTE:

This figure is to be read in conjunction with the accompanying Golder Associates Ltd. report No. 11-1127-0134

### REFERENCE

Base data provided by DMTI Spatial Inc, 2006, NRVIS, 2004, used under license. Projection: Transverse Mercator Datum: NAD 83 Coordinate System: UTM Zone 18



PROJECT

### RICHMOND VILLAGE GROUNDWATER VULNERABILITY

TITLE

### KINGS PARK UPPER AQUIFER PARTICLE TRACES



PROJECT No. 11-1127-0134				
DESIGN	NB	Jan. 2013		
GIS	ABD	Jan. 2013		
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SCALE AS SHOWN REV. 2 FIGURE: C2



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### NOTE:

This figure is to be read in conjunction with the accompanying Golder Associates Ltd. report No.11-1127-0134

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PROJECT

### RICHMOND VILLAGE GROUNDWATER VULNERABILITY

MUNSTER HAMLET AND KINGS PARK COMBINED LOWER AQUIFER PARTICAL TRACES



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SCALE AS SHOWN REV. 2



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<del>\$</del>	WATER SUPPLY WELL
	Roadway
-+	Railway
•	Utility Line
	Watercourse, Permanent
	Watercourse, Intermittent
	Building
	Water Area, Permanent
<u></u>	Wetland, Permanent
	Munster Hamlet Boundary
	Upper Aquifer Wellhead Protection Area
Uppe	r Aquifer Intrinsic Vulnerability
	High
	Medium

### NOTE:

PROJECT

TITLE

This figure is to be read in conjunction with the accompanying Golder Associates Ltd. report No.11-1127-0134

### REFERENCE

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RICHMOND VILLAGE GROUNDWATER VULNERABILITY

### MUNSTER HAMLET - UPPER AQUIFER INTRINSIC VULNERABILITY

	PROJECT No. 11-1127-0134		SCALE AS SHOWN	REV. 2	
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	REVIEW				



WATER SUPPLY WELL
Roadway
Railway
Utility Line
Watercourse, Permanent
Watercourse, Intermittent
Building
Water Area, Permanent
Wetland, Permanent
Upper Aquifer Wellhead Protection Area
Village of Richmond Boundary
Upper Aquifer Intrinsic Vulnerability
High

Medium

### NOTE:

This figure is to be read in conjunction with the accompanying Golder Associates Ltd. report No. 11-1127-0134

### REFERENCE

Base data provided by DMTI Spatial Inc, 2006, NRVIS, 2004, used under license. Projection: Transverse Mercator Datum: NAD 83 Coordinate System: UTM Zone 18



PROJECT

TITLE

### RICHMOND VILLAGE GROUNDWATER VULNERABILITY

### KINGS PARK - UPPER AQUIFER INTRINSIC VULNERABILITY



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GIS	ABD	Jan. 2013		
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REVIEW				

FIGURE: C5



High Medium

Low

♦ WATER SUPPLY WELL	÷			
Roadway				
H Railway				
Utility Line	•			
Watercourse, Permanent				
Watercourse, Intermittent				
Building				
Water Area, Permanent				
Wetland, Permanent	$d\mu^{-3}$			
Upper and Lower Aquifer Intrinsic Vulnerability				

### NOTE:

This figure is to be read in conjunction with the accompanying Golder Associates Ltd. report No. 11-1127-0134

### REFERENCE

Base data provided by DMTI Spatial Inc, 2006, NRVIS, 2004, used under license. Projection: Transverse Mercator Datum: NAD 83 Coordinate System: UTM Zone 18



PROJECT

### RICHMOND VILLAGE GROUNDWATER VULNERABILITY

## MUNSTER HAMLET AND KINGS PARK COMBINED AQUIFER INTRINSIC VULNERABILITY



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SCALE AS SHOWN REV. 2









### NOTE

THIS FIGURE IS TO BE READ IN CONJUNCTION WITH THE ACCOMPANYING GOLDER ASSOCIATES LIMITED REPORT NO. 11-1127-0134

### REFERENCE

BASE DATA - MNR NRVIS, OBTAINED 2008, CANMAP V2008.4 PRODUCED BY GOLDER ASSOCIATES LTD UNDER LICENCE FROM ONTARIO MINISTRY OF NATURAL RESOURCES, © QUEENS PRINTER 2008 PROJECTION: TRANSVERSE MERCATOR DATUM: NAD 83 COORDINATE SYSTEM: UTM ZONE 18

1,200	0	1,200
SCALE 1:60,000		METRES

PROJECT

TITLE

### RICHMOND VILLAGE GROUNDWATER VULNERABILITY

### KINGS PARK - UPPER AND LOWER AQUIFER COMBINED INTRINSIC VULNERABILITY SCORE

	PROJECT No. 11-1127-0134			SCALE AS SHOWN	REV. 2	
Golder Associates Ottawa, Ontario	DESIGN	NB	Jan. 2013			
	GIS	ABD	Jan. 2013			
	CHECK				60	
	REVIEW					

![](_page_84_Figure_0.jpeg)

¢	WATER SUPPLY WELL					
	Roadway					
	Railway					
	Utility Line					
	Watercourse, Permanent					
	Watercourse, Intermittent					
	Building					
	Water Area, Permanent					
() sk.	Wetland, Permanent					
	Munster Hamlet Boundary					
Vulne	Vulnerability Score					
	Low Uncertainty					

High Uncertainty

### NOTE:

TITLE

This figure is to be read in conjunction with the accompanying Golder Associates Ltd. report No. 11-1127-0134

### REFERENCE

Base data provided by DMTI Spatial Inc, 2006, NRVIS, 2004, used under license. Projection: Transverse Mercator Datum: NAD 83 Coordinate System: UTM Zone 18

![](_page_84_Figure_8.jpeg)

RICHMOND VILLAGE GROUNDWATER VULNERABILITY

### MUNSTER HAMLET - UPPER AQUIFER VULNERABILITY SCORE UNCERTAINTY

	PROJECT No. 11-1127-0134			SCALE AS SHOWN	REV.
	DESIGN	NB	Jan. 2013		
Golder	GIS	ABD	Jan. 2013	FIGURE:	C9
Associates	CHECK				
Ottawa, Ontario	REVIEW				

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![](_page_85_Figure_0.jpeg)

<del>\$</del>	WATER SUPPLY WELL						
	Roadway						
	Railway						
•	Utility Line						
	Watercourse, Permanent						
	Watercourse, Intermittent						
	Building						
	Water Area, Permanent						
	Wetland, Permanent						
Vulne	Vulnerability Score						
	High Uncertainty						

Low Uncertainty

### NOTE:

This figure is to be read in conjunction with the accompanying Golder Associates Ltd. report No11-1127-0134

### REFERENCE

Base data provided by DMTI Spatial Inc, 2006, NRVIS, 2004, used under license. Projection: Transverse Mercator Datum: NAD 83 Coordinate System: UTM Zone 18

![](_page_85_Figure_9.jpeg)

METRES

PROJECT

TITLE

### RICHMOND VILLAGE GROUNDWATER VULNERABILITY

### KINGS PARK - UPPER AQUIFER UNCERTAINTY

		PROJECT	No. 11-	1127-0134	SCALE AS SHOWN	REV. 2
	Golder Associates Ottawa, Ontario	DESIGN	NB	Jan. 2013	FIGURE:	C10
		GIS	ABD	Jan. 2013		
		CHECK				
		REVIEW				

![](_page_86_Figure_0.jpeg)

![](_page_86_Figure_3.jpeg)

### Vulnerability Score Low Uncertainty High Uncertainty

### NOTE:

This figure is to be read in conjunction with the accompanying Golder Associates Ltd. report No. 11-1127-0134

### REFERENCE

Base data provided by DMTI Spatial Inc, 2006, NRVIS, 2004, used under license. Projection: Transverse Mercator Datum: NAD 83 Coordinate System: UTM Zone 18

![](_page_86_Figure_9.jpeg)

PROJECT

TITLE

### RICHMOND VILLAGE GROUNDWATER VULNERABILITY

MUNSTER HAMLET AND KINGS PARK COMBINED LOWER AQUIFER VULNERABILITY UNCERTAINTY

![](_page_86_Picture_13.jpeg)

PROJECT	PROJECT No. 11-1127-0134 SCALE AS SHOWN			REV. 2		
DESIGN	NB	Jan. 2013				
GIS	ABD	Jan. 2013				
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REVIEW						

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![](_page_87_Picture_5.jpeg)